

# **Fusion partielle, granites et migmatites la dynamique des croûtes chaudes archéennes**

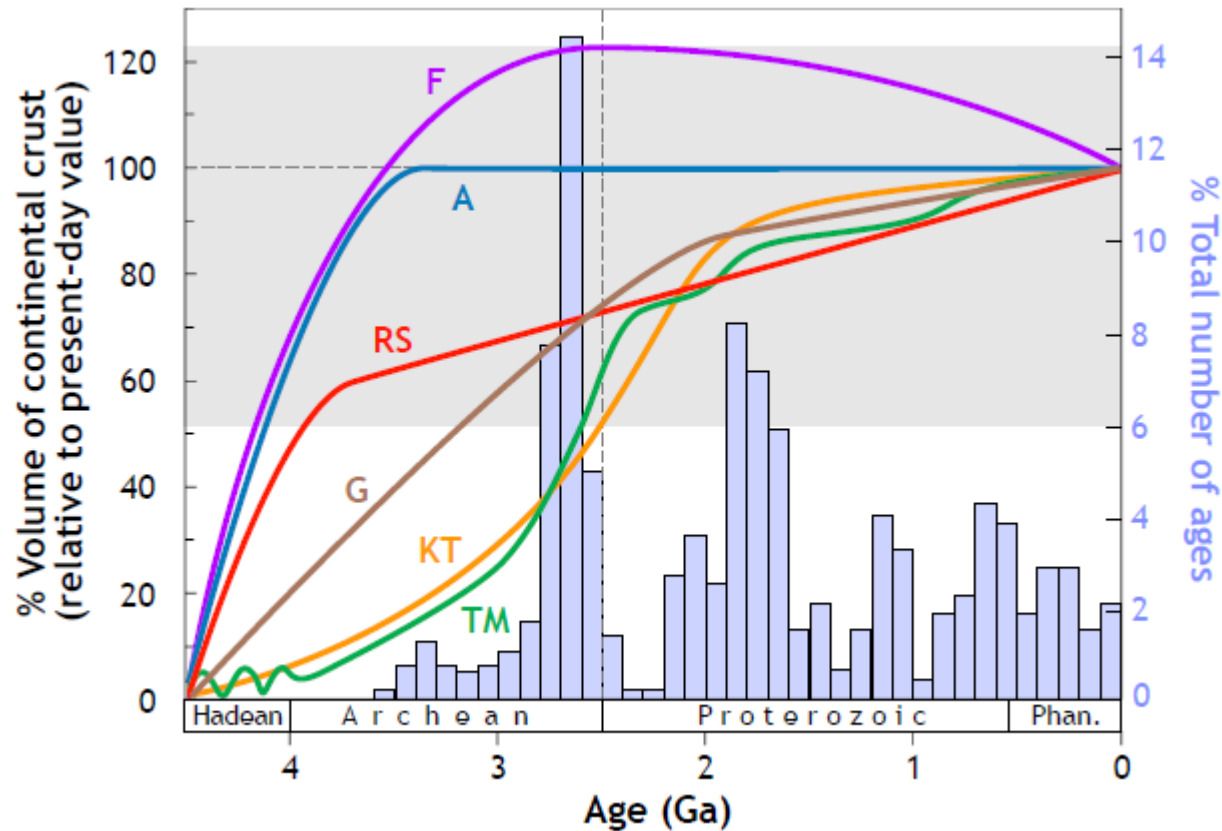
**Jean-François Moyen**

Avec des idées empruntées à  
H. Martin, O. Vanderhaeghe, O. Laurent

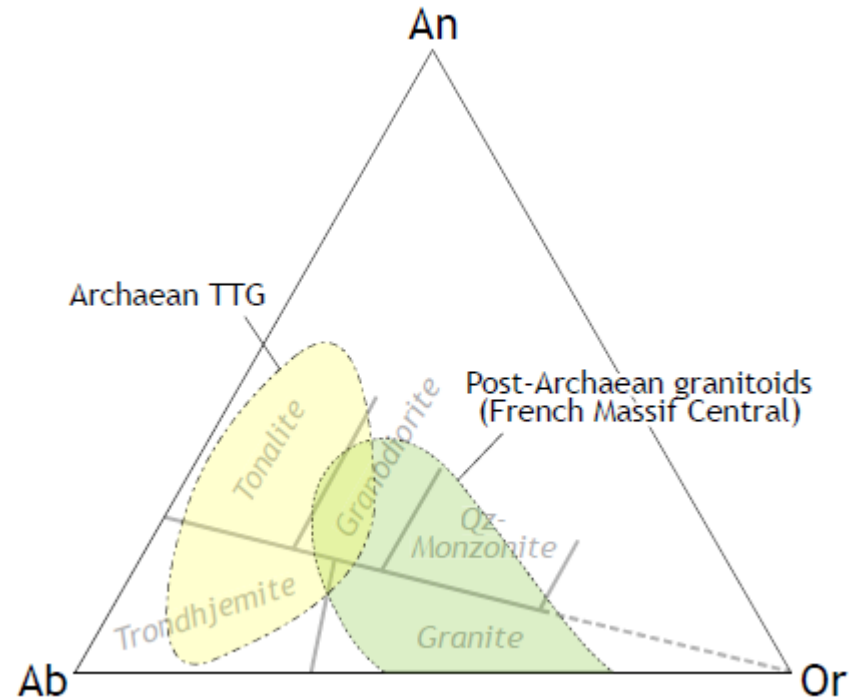
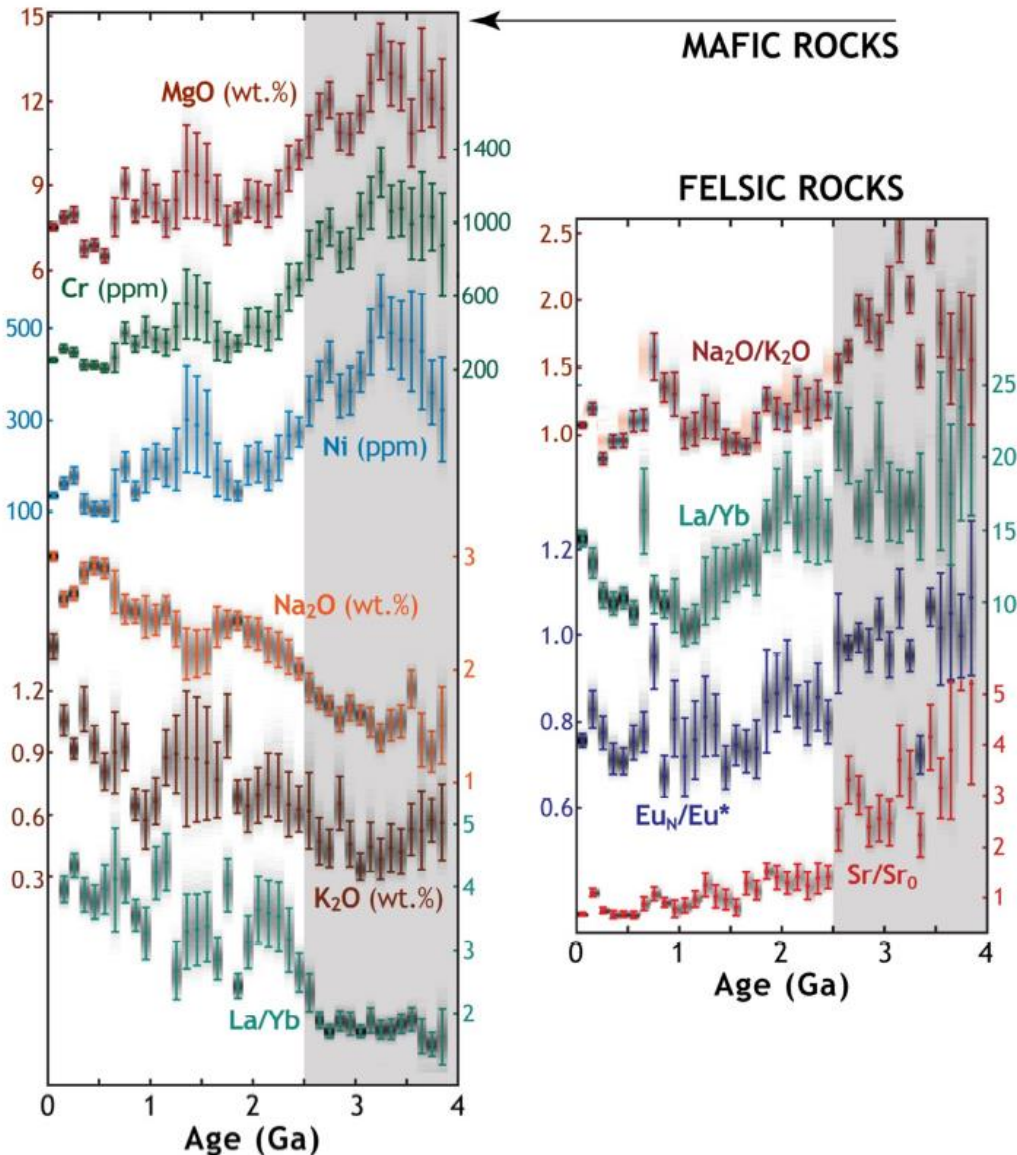
# De quoi ça parle?

- Les granitoïdes et la construction de la croûte archéenne
- Que nous dit la géochimie des granitoïdes?
- Les granitoïdes comme marqueurs de la dynamique de la croûte
- Un peu de géodynamique

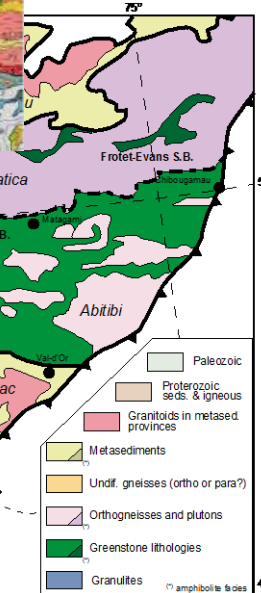
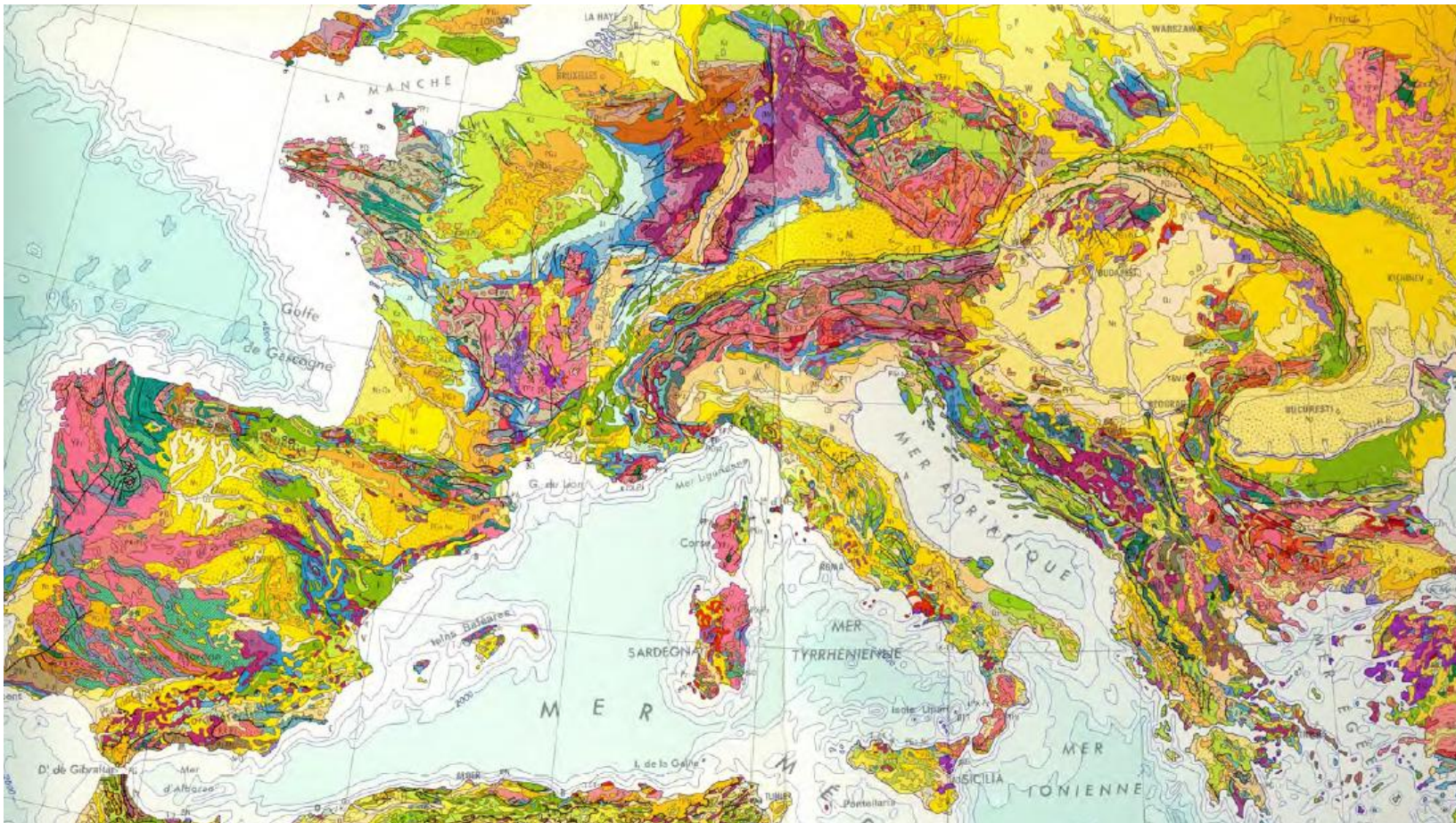
# La croûte continentale est surtout Archéenne



# La CC archéenne est différente de la CC moderne

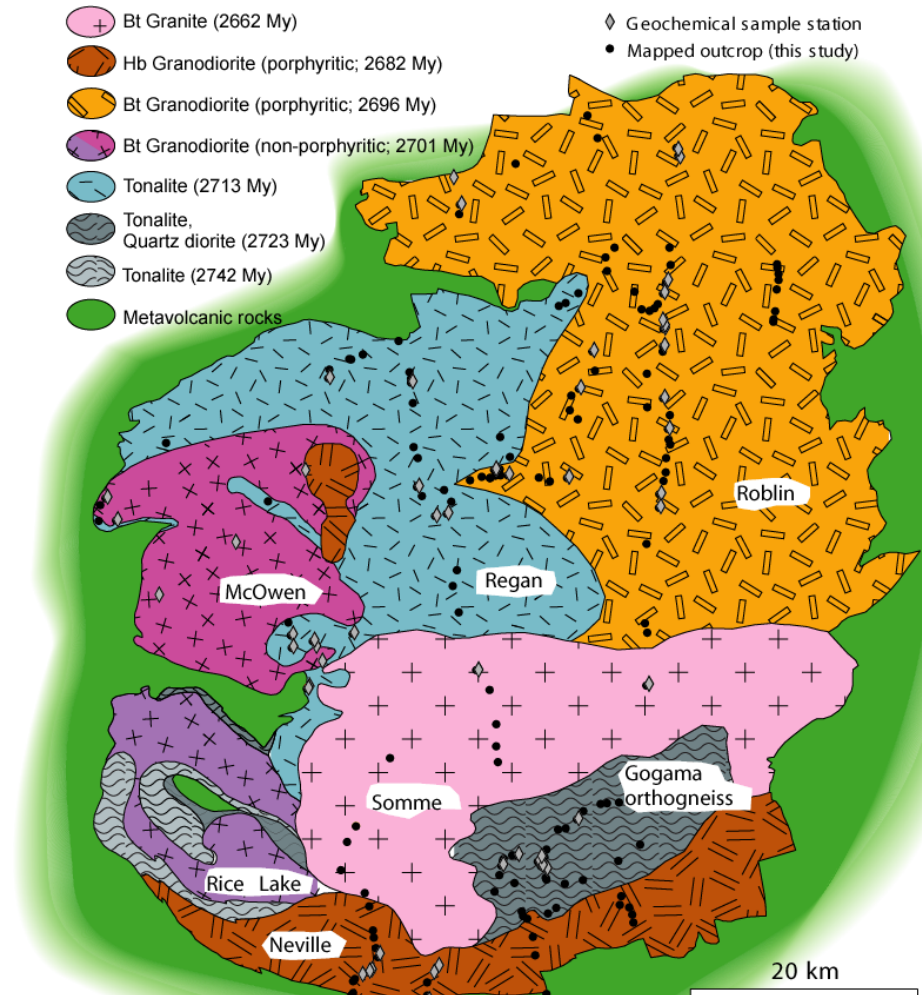
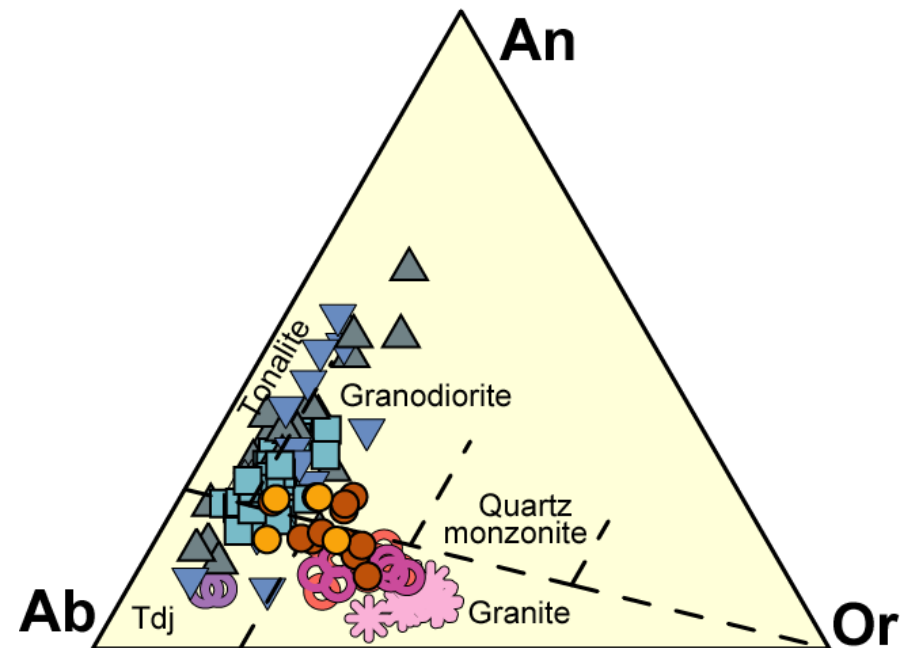
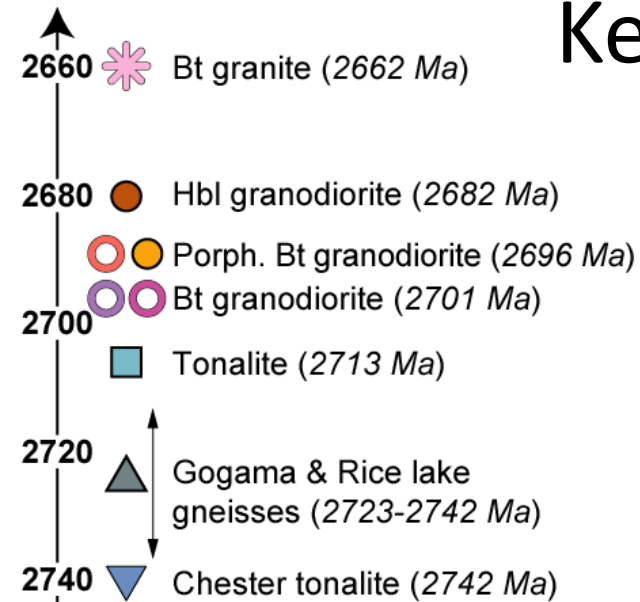


*... et les granites aussi !*



Proto-croûte  
Vs.  
Croûte felsique évoluée

# Evolution progressive de la crt : Kenogamissi complex



# Archaean geology: the basic question

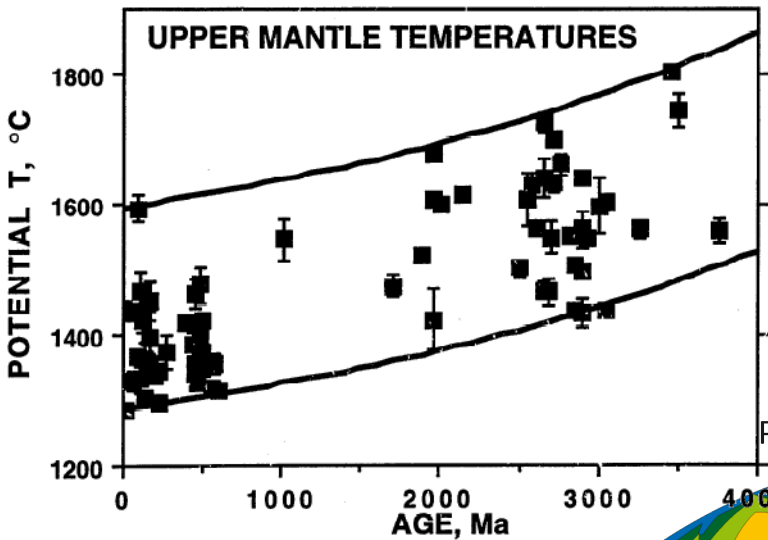
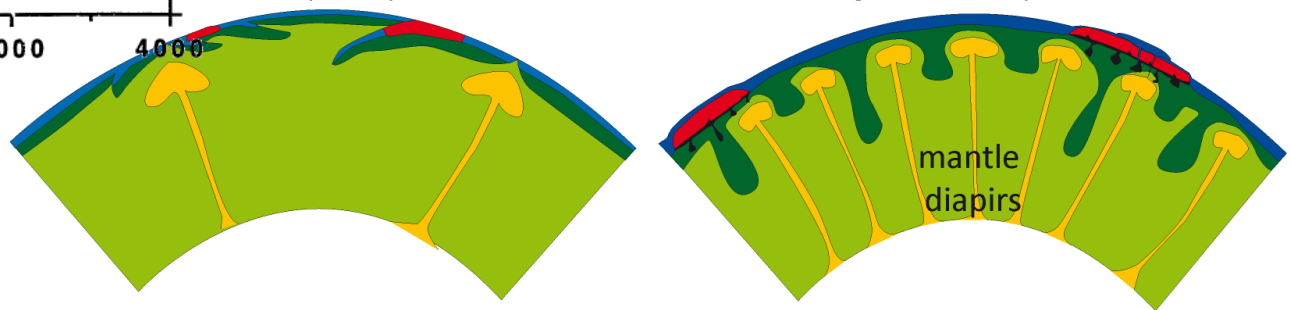


Plate tectonics or not in the hotter Archaean mantle ?

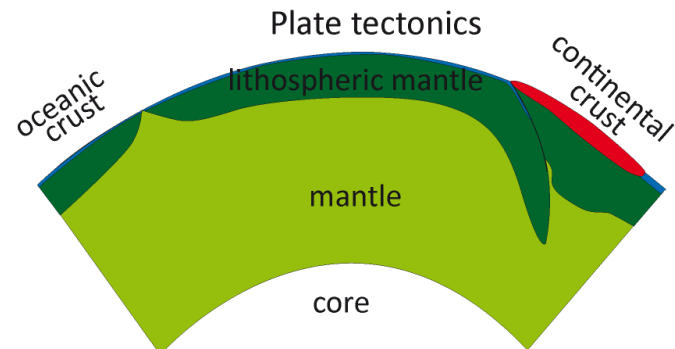
Archaean Geodynamics

Platelets-protoplates

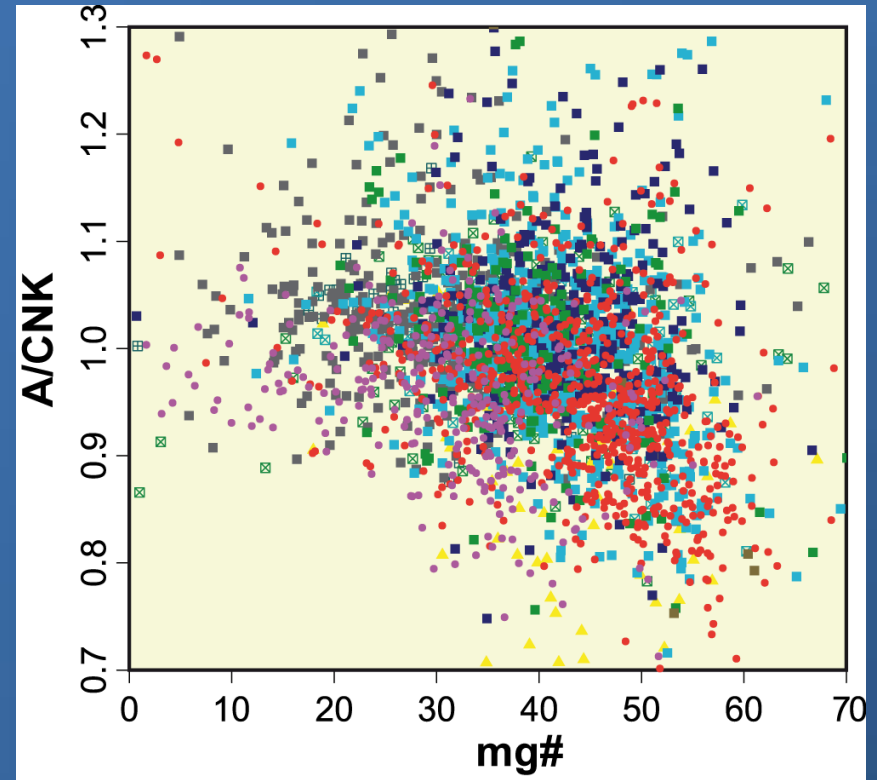
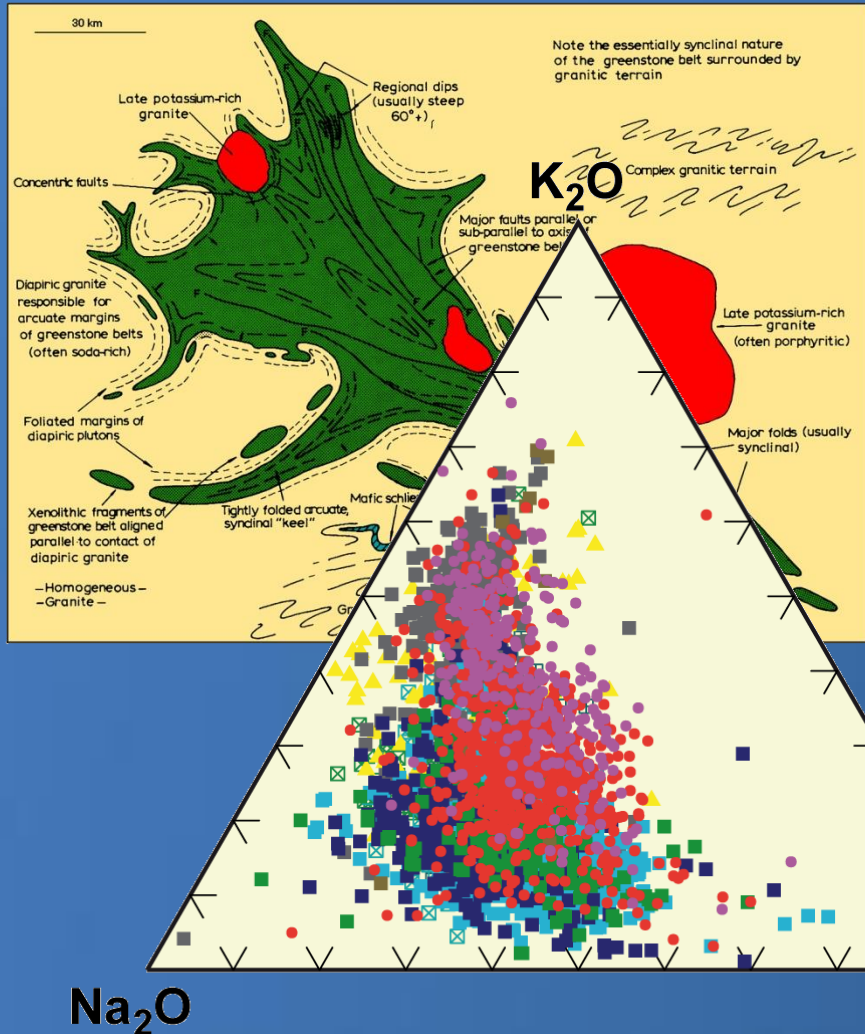
Stagnant lithospheric lid



Phanerozoic Geodynamics



# Géochimie des granites Archéens



$SiO_2 > 60\%$   
 $N \approx 3000$

# A common approach: « tectonic discrimination »

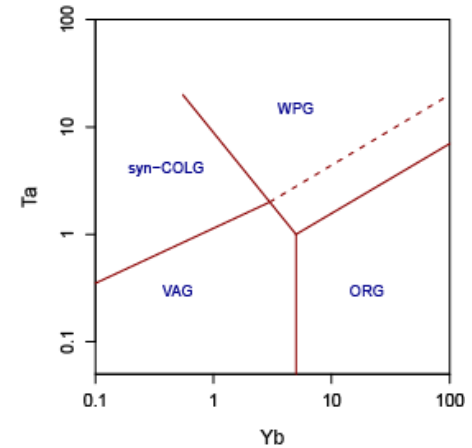
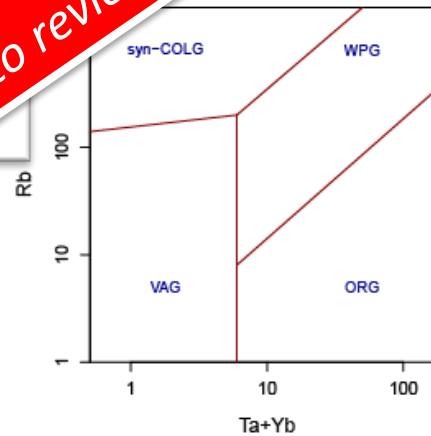
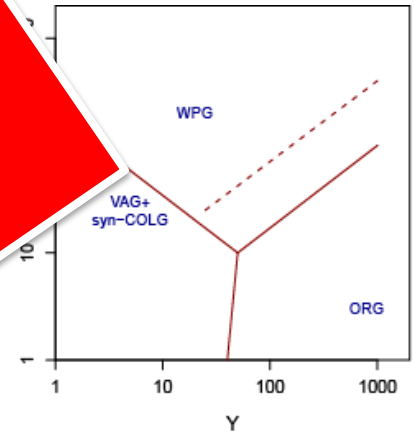
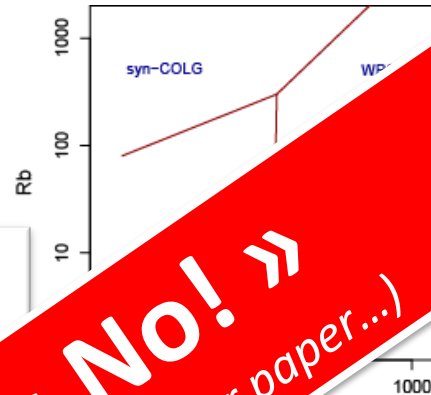
## Trace Element Discrimination Diagrams for the Tectonic Interpretation of Granitic Rocks

by JULIAN A. PEARCE\*, NIGEL B. W. HARRIS  
ANDREW G. TINDLE

Department of Earth Sciences, The Open University, Milton Keynes  
England

(Received 2 February 1984; in revised form 15 May 1984)

**Just say « No! »**  
(or don't ask me to review your paper...)



# What's wrong with geotectonic (« discriminant ») diagrams for (Archaean) granites?

1. They assume a finite, well-defined set of possible tectonic sites (e.g. no post-orogenic setting ?!)
2. They assume we know possible Archaean tectonic sites (!)
3. They do not tell anything you don't know before...
4. They do not work anyway (at least for complex settings)

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### *Post-collision granites*

This group represents a major problem in all tectonic-geochemical classifications of granites. The probable reason is that, unlike granites from other settings, post-collision granites cannot be explained in terms of a single, well-defined mantle or crustal source. It has

Thus, at present, geological criteria and the zonation of intrusions are the best indicators of a post-orogenic setting. However, if the setting can be defined as 'post-collision' on *a priori*

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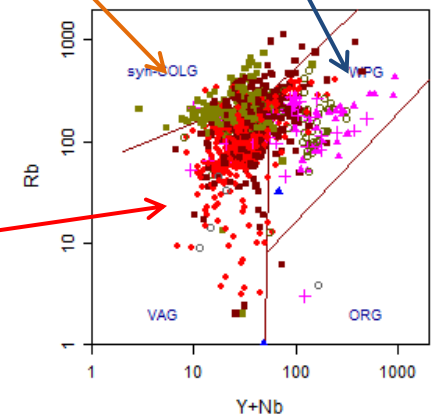
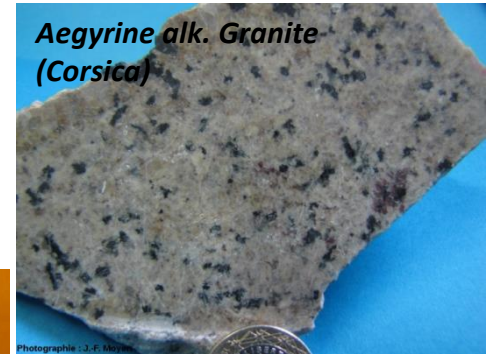
### *Age*

Since the diagrams are based on analyses of Phanerozoic rocks they cannot automatically be applied to rocks of Precambrian age. In detail, the main differences between magma genesis in the Archaean and at the present day are likely to be:

- (i) The mantle evolution line in Fig. 5a will be displaced to higher Rb values, and very slightly higher (Y + Nb) values.
  - (ii) The degree of partial melting *may* be greater.
  - (iii) The mantle may be less fractionated than at the present-day so that generation of magmas from very enriched and very depleted sources would be less common.
  - (iv) Crustal melting *may* be more important in magma genesis in all environments.
- The net effect on the diagrams should be to blur the syn-COLG/VAG discrimination and displace the (VAG + syn-COLG)/(WPG + ORG) boundary to lower values of Y, Yb, (Y + Nb) and (Yb + Ta). However, the precise effect cannot be known until further studies have been made on granites of varying ages.

# What's wrong with geotectonic (« discriminant ») diagrams for (Archaean) granites?

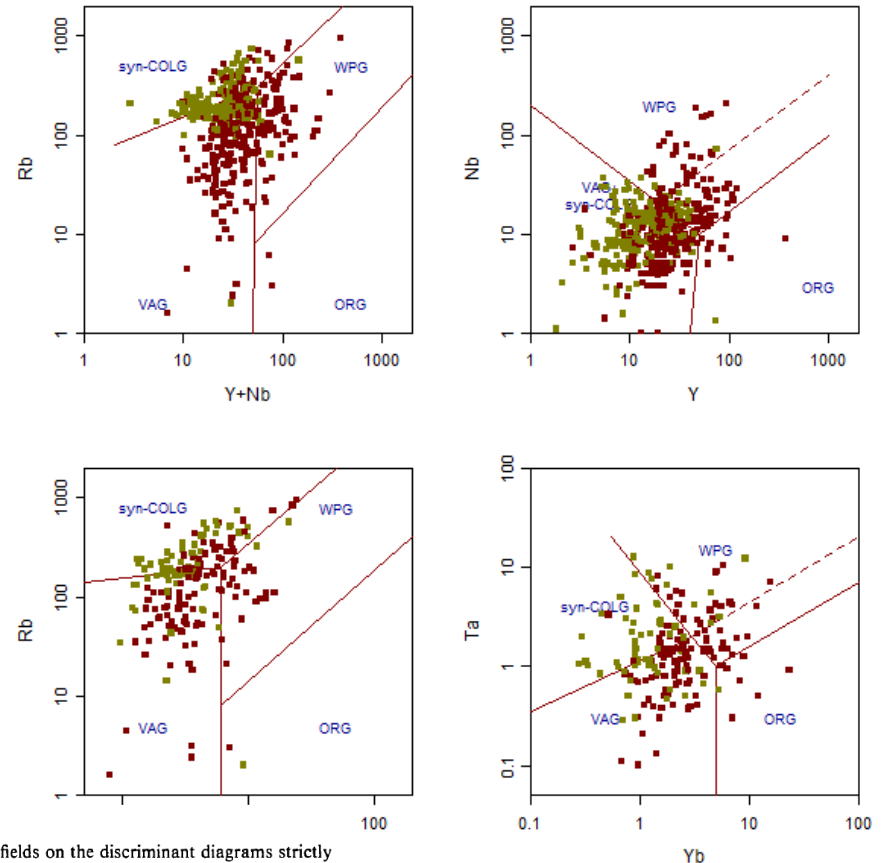
1. They assume a finite, well-defined set of possible tectonic sites (e.g. no post-orogenic setting ?!)
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Granite tectonic discrimination – Pearce et al. (1984)



## Complex tectonic regimes

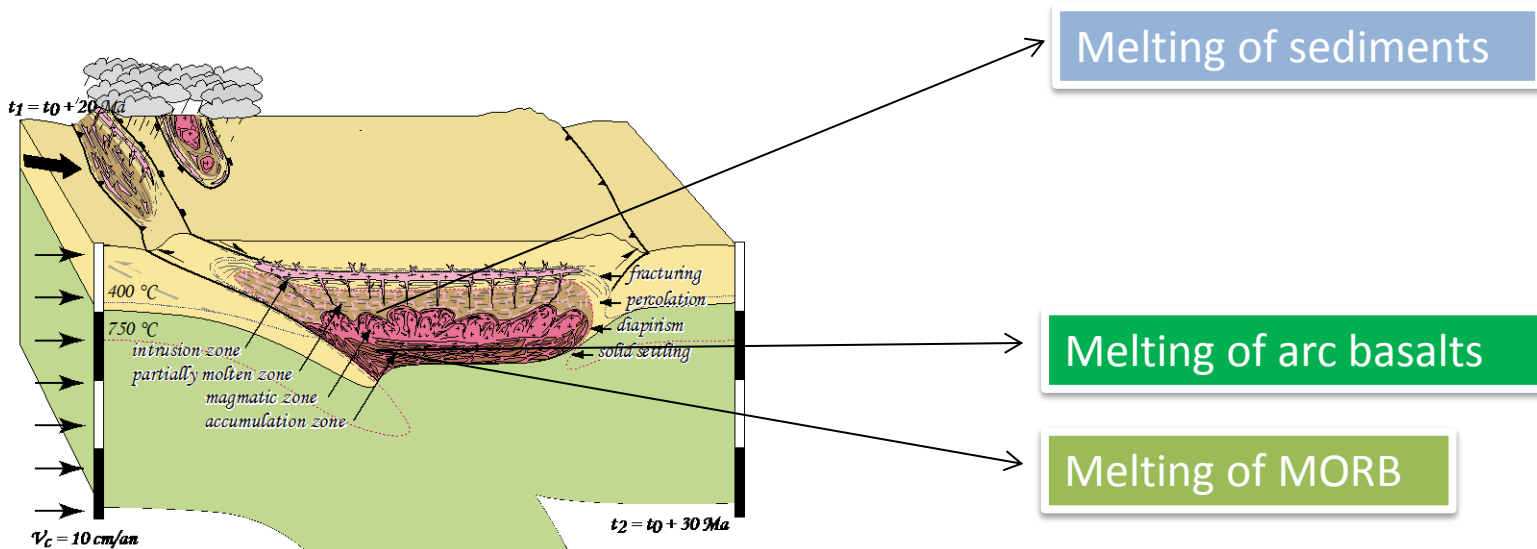
As we emphasized in the previous section, the fields on the discriminant diagrams strictly reflect source regions (and melting and crystallization histories) rather than tectonic regimes. For syn-COL, VA, 'normal' OR and WP granites source regions and tectonic regimes show a clear correlation. However, post-COL granites can plot in the syn-COLG, VAG or WPG fields and OR granites can plot in the VAG or ORG fields according to their precise setting. Thus, the discrimination is not perfect between *all* groups and interpretations must be made with this caveat in mind.

A further cause for caution concerns the complex nature, not only of collision settings, but of many apparently simple tectonic regimes. We have already pointed out, for example, that

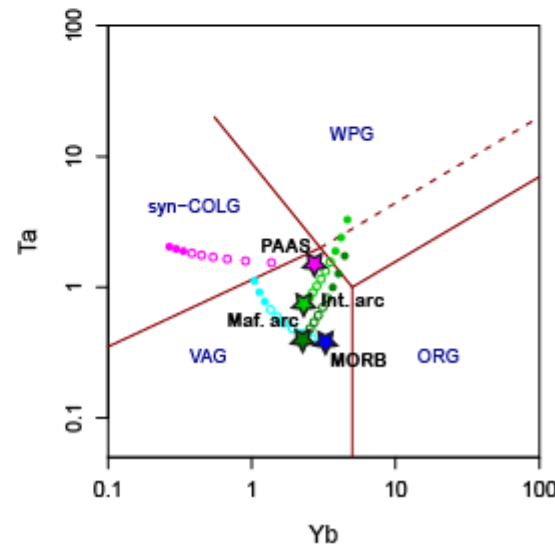
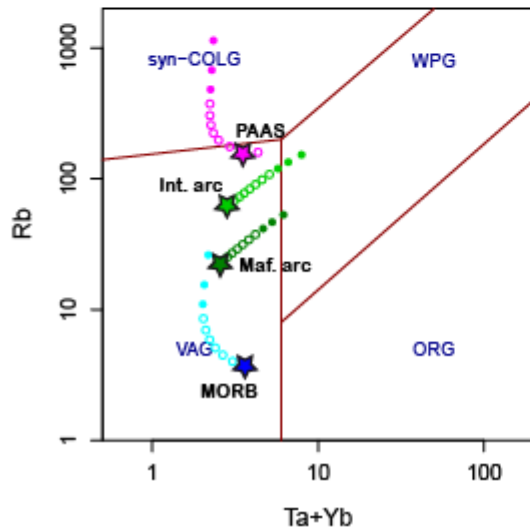
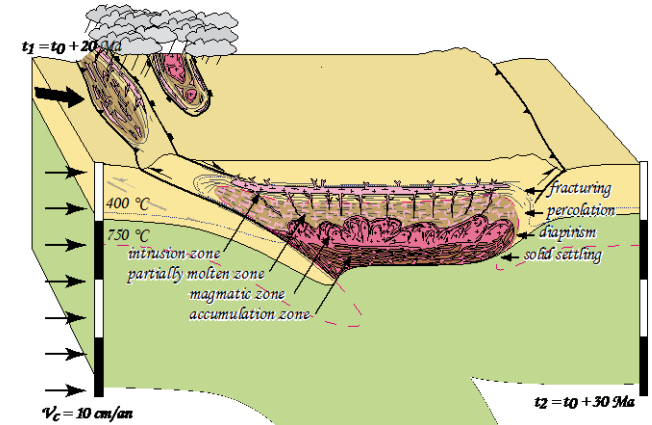
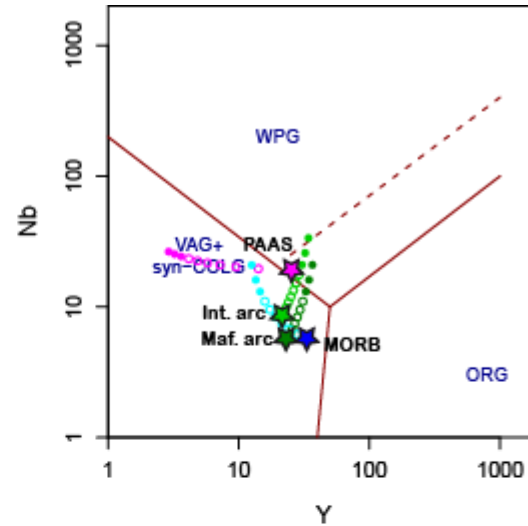
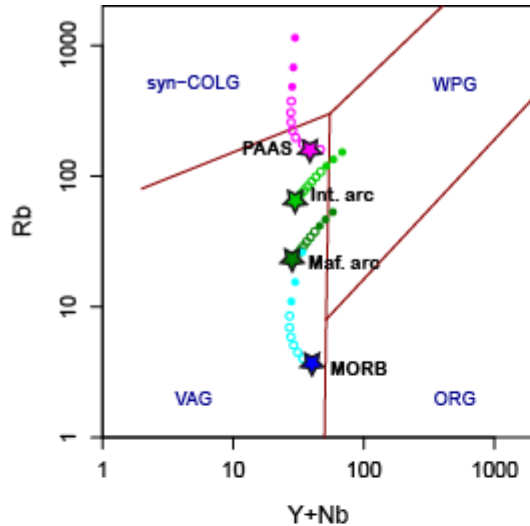
# So, what can we do ?

- Remember that

$$C_1 = \frac{C_0}{F + D(1 - F)}$$

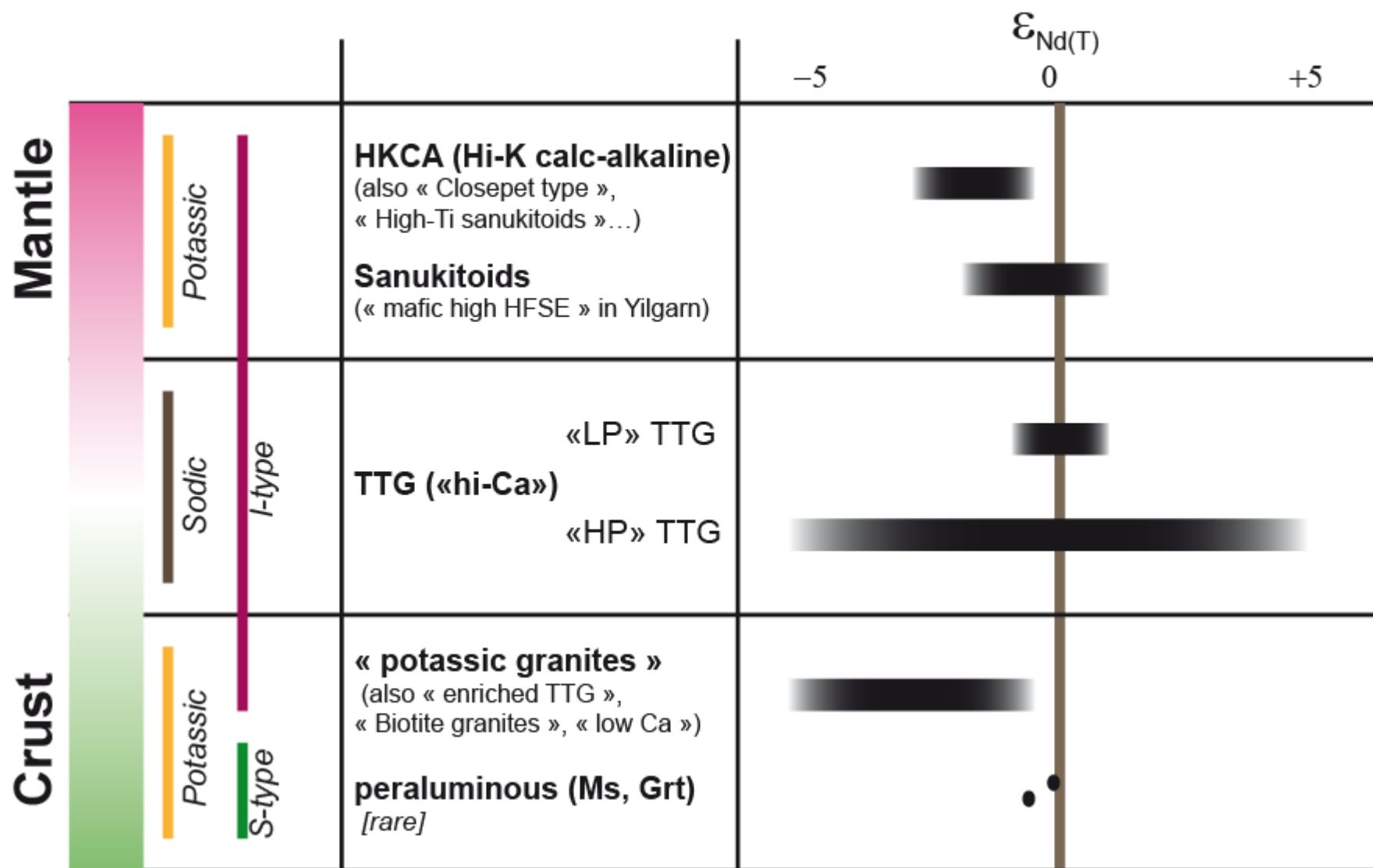


# Apparently different tectonic sites!

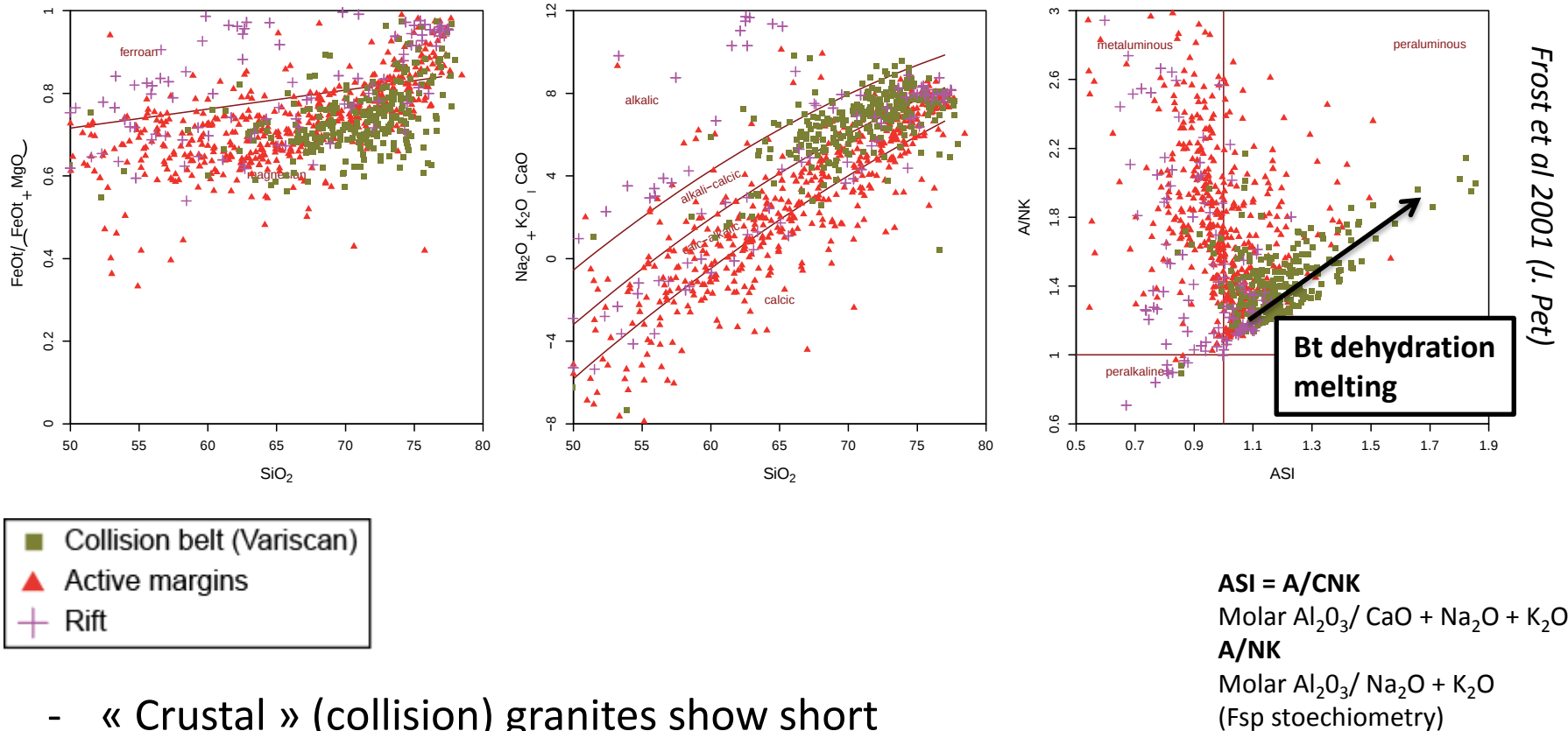


*But, of course,  
everything happens in  
the same place  
(= same setting...)*

# Types of Archaean granites



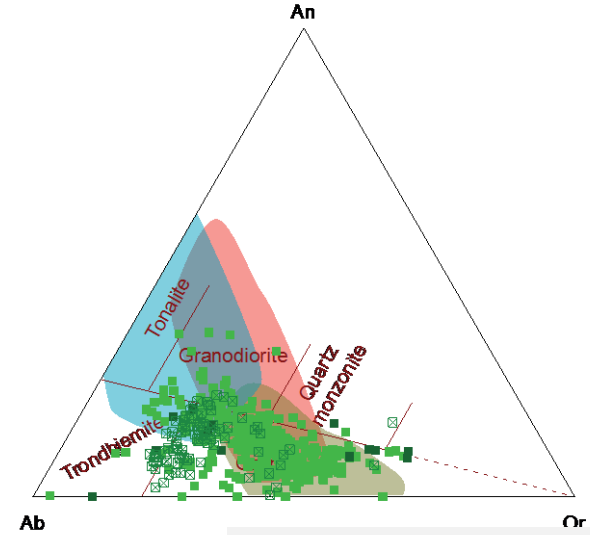
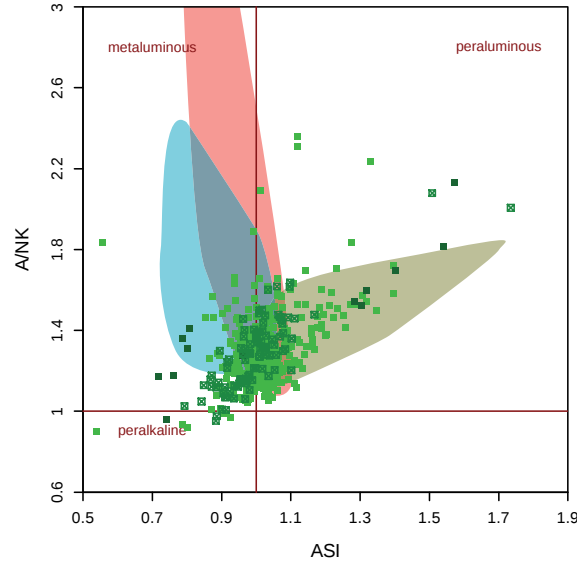
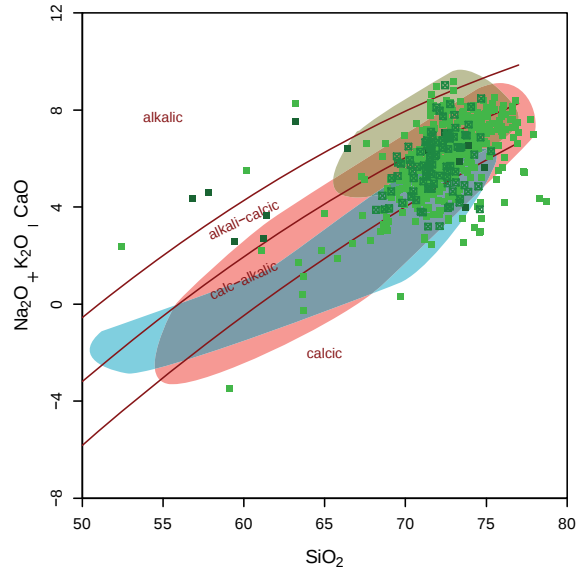
# Modern (phanerozoic) granites



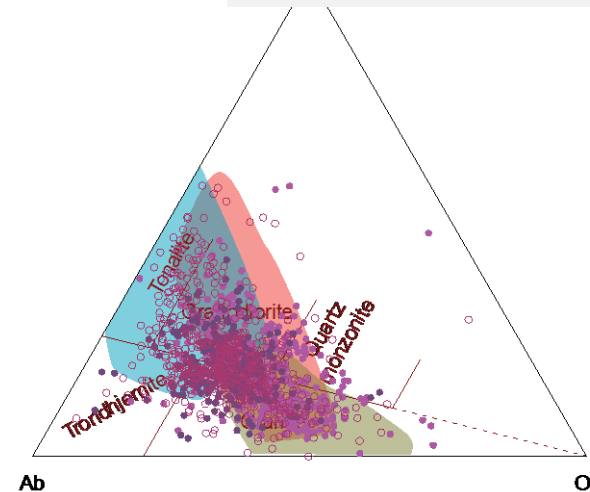
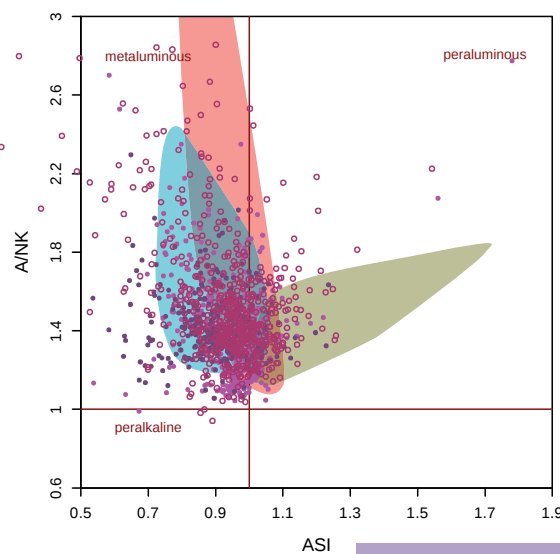
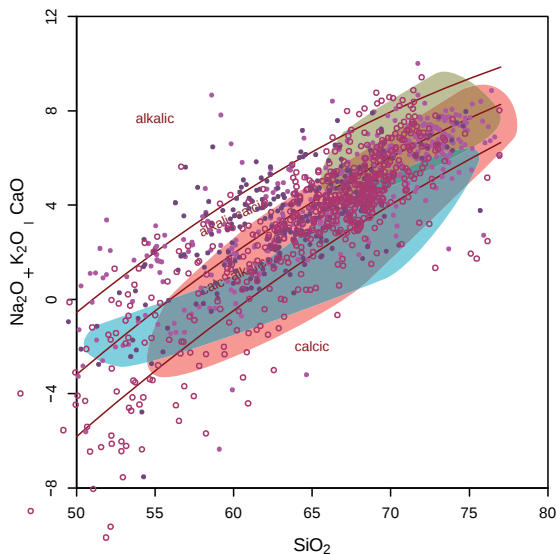
- « Crustal » (collision) granites show short trends (no intermediate/mafic rocks associated)
- Very distinctive behaviour in A/CNK vs. A/NK (cst Ca)

→ Fundamental difference in petro. behaviour between both types!

# Some Archaean granites are comparable to modern



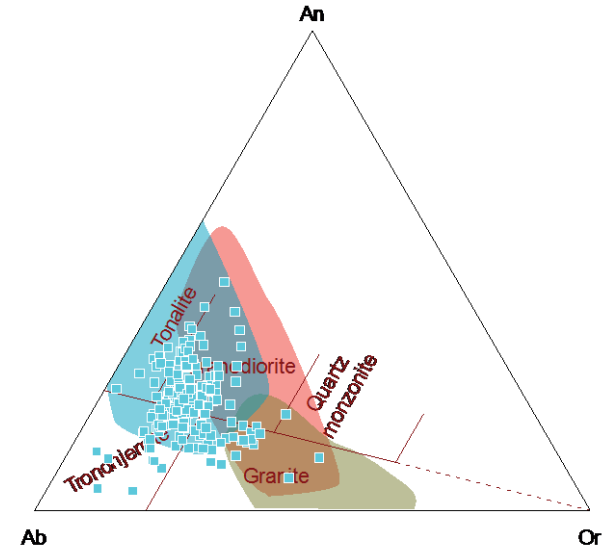
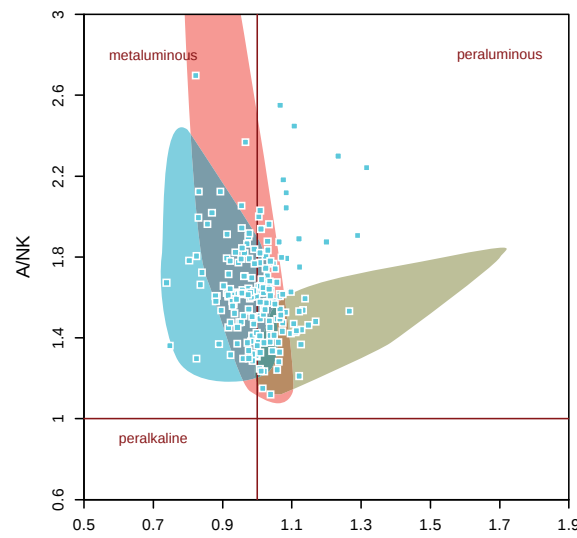
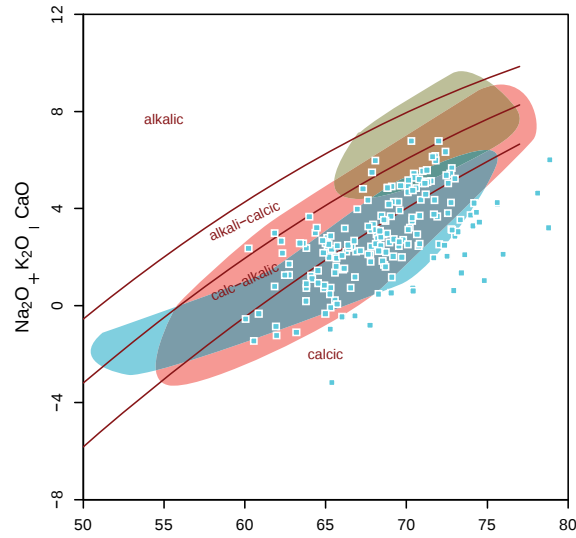
Fields:  
Brown- collision  
Red- Active margin/post coll  
Blue- Oceanic plateau/ridge



Mantle (HKCA, sanukitoids) and post-collision

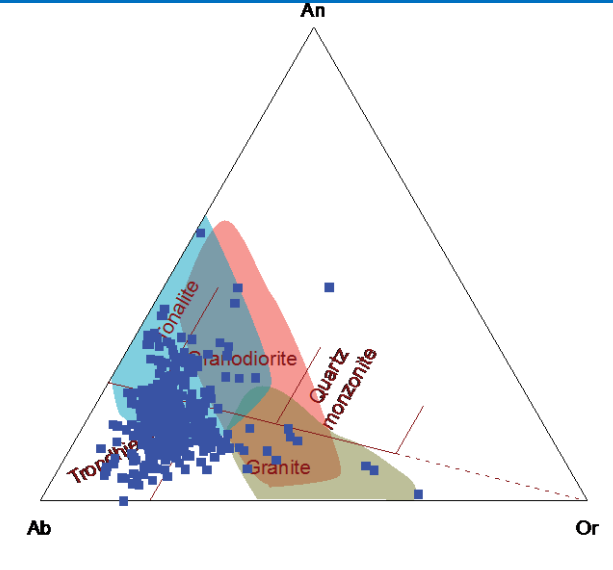
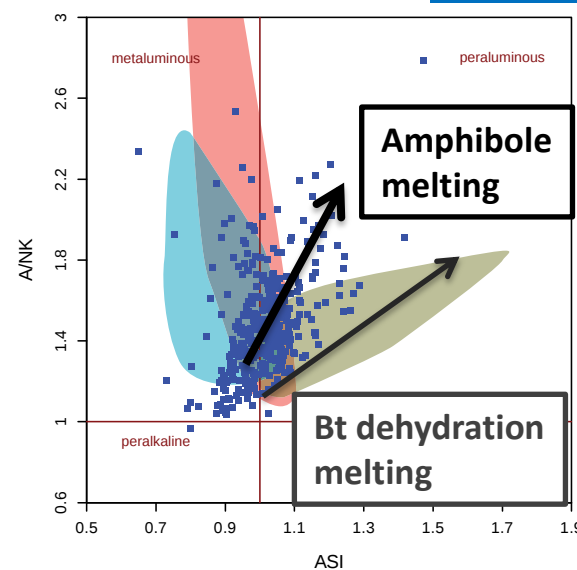
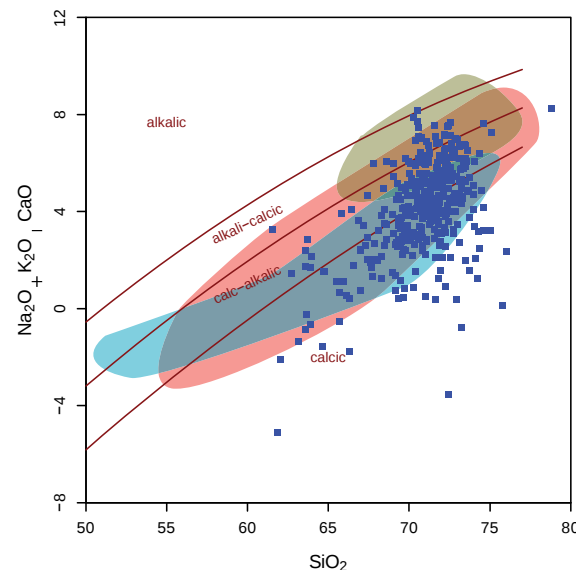
Crustal granites and collisions

# No real equivalent for TTGs



« low pressure » (ca. 10 kbar evolution) TTGs and oceanic plateaus/ridges

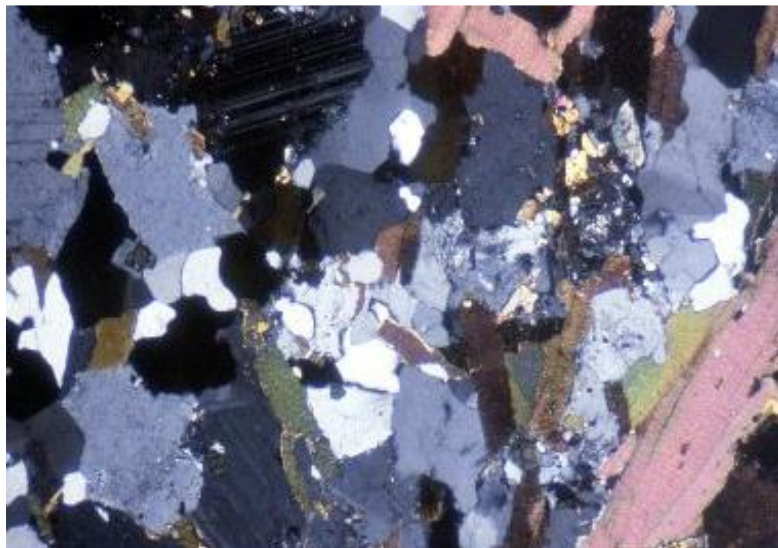
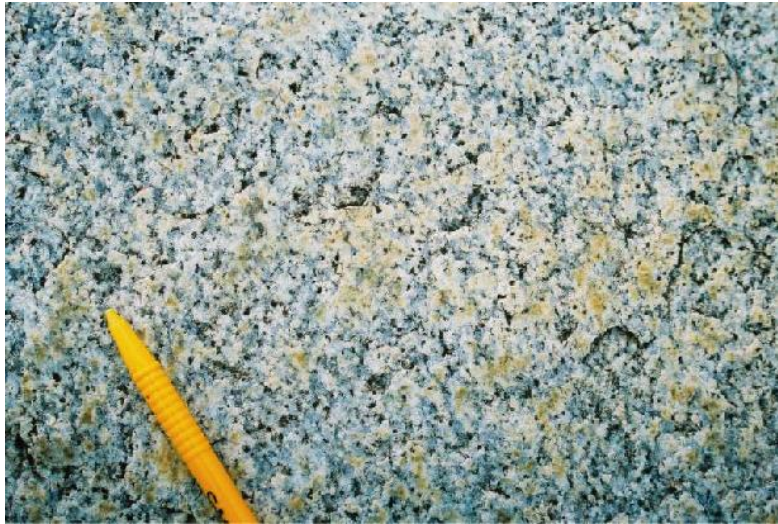
« high pressure » TTGs (20 kbar) and ???



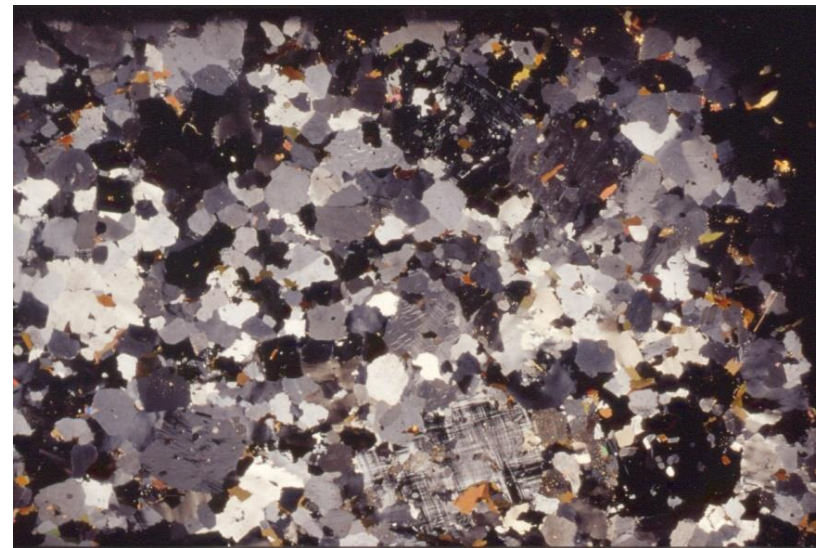
# « Crust » vs. « Mantle » granites

| Mantle origin                                            | Crust origin                                             |
|----------------------------------------------------------|----------------------------------------------------------|
| Comagmatic microgranular mafic enclaves (MME) ubiquitous | MME rare                                                 |
| Generally Hornblende bearing                             | Hornblende rare                                          |
| Generally potassic                                       | Potassic or sodic                                        |
| Often K-feldspar porphyritic                             | Variable, more frequently equigranular                   |
| Syn to post tectonic                                     | Variable from pre to post tectonic                       |
| Little zircon inheritance                                | Significant zircon inheritance (old cores etc.)          |
| Relatively low A/CNK<br>negative slope in Shand diagram  | Relatively high A/CNK<br>positive slope in Shand diagram |
| Relatively high mg#                                      | Relatively low mg#                                       |

# Archaean C-type granites

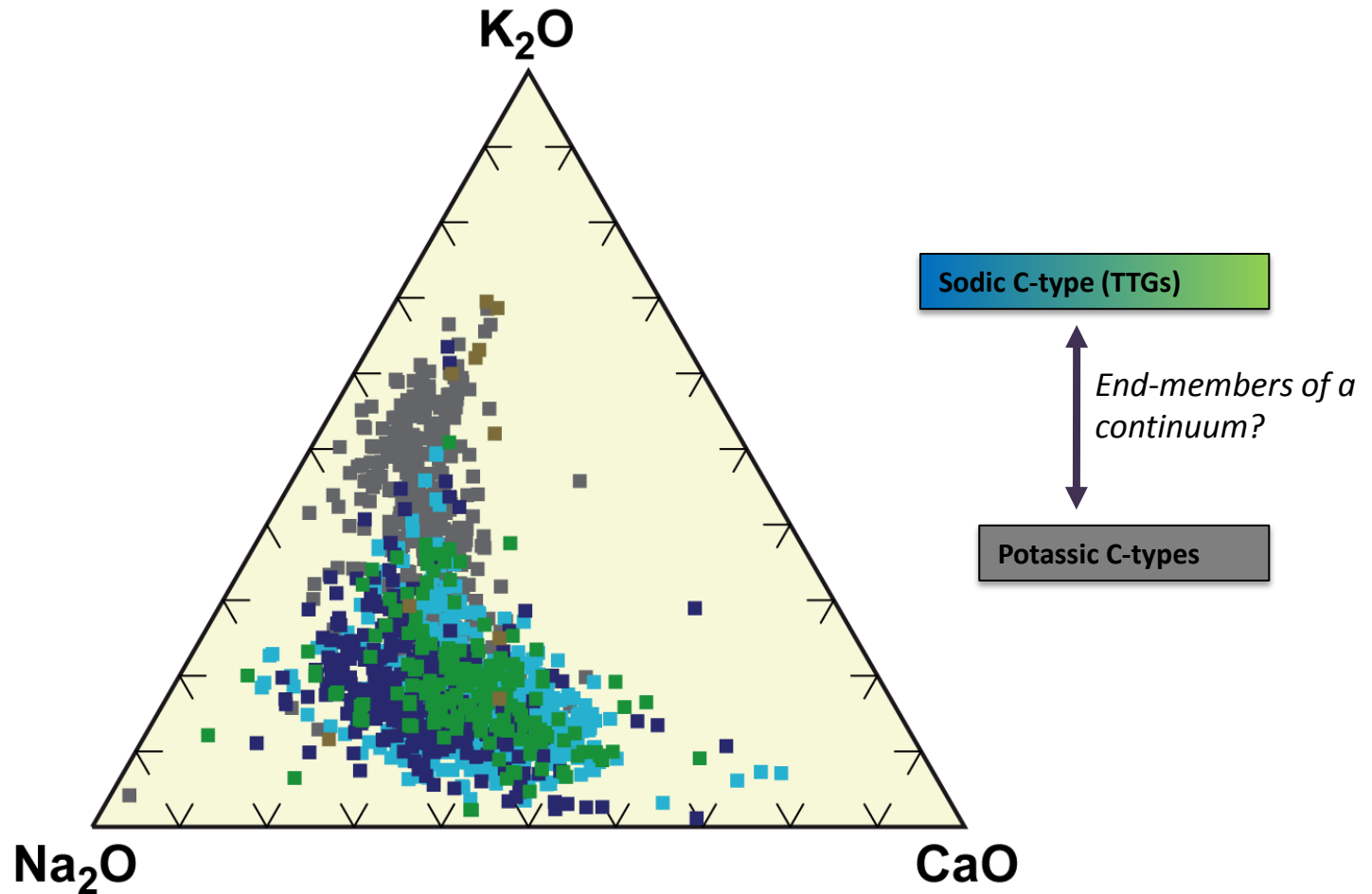


Pg + Qz + Bt ( $\pm$  amp, Ep)



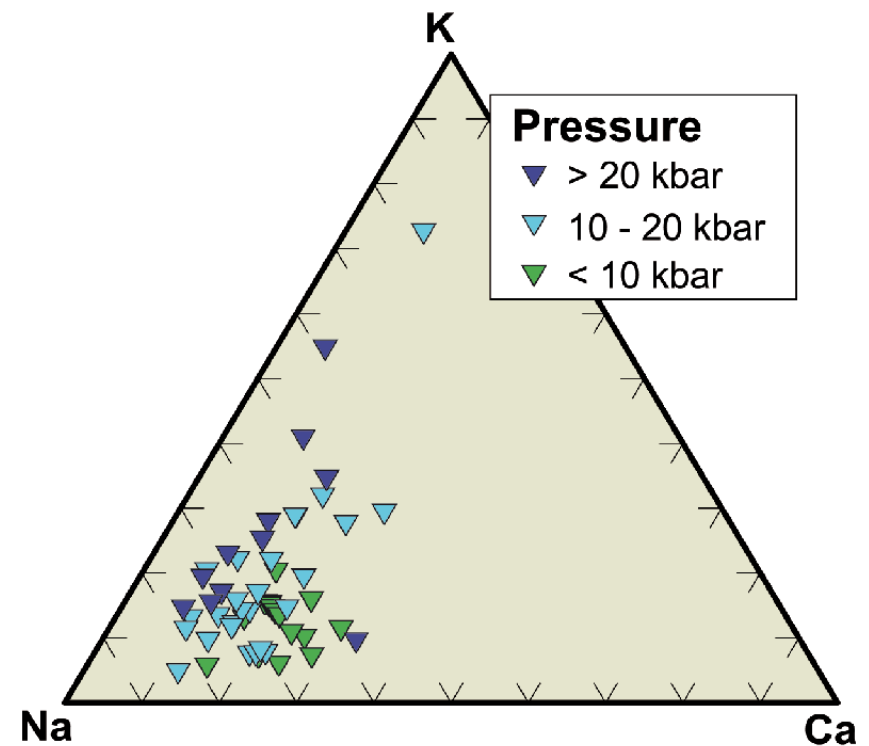
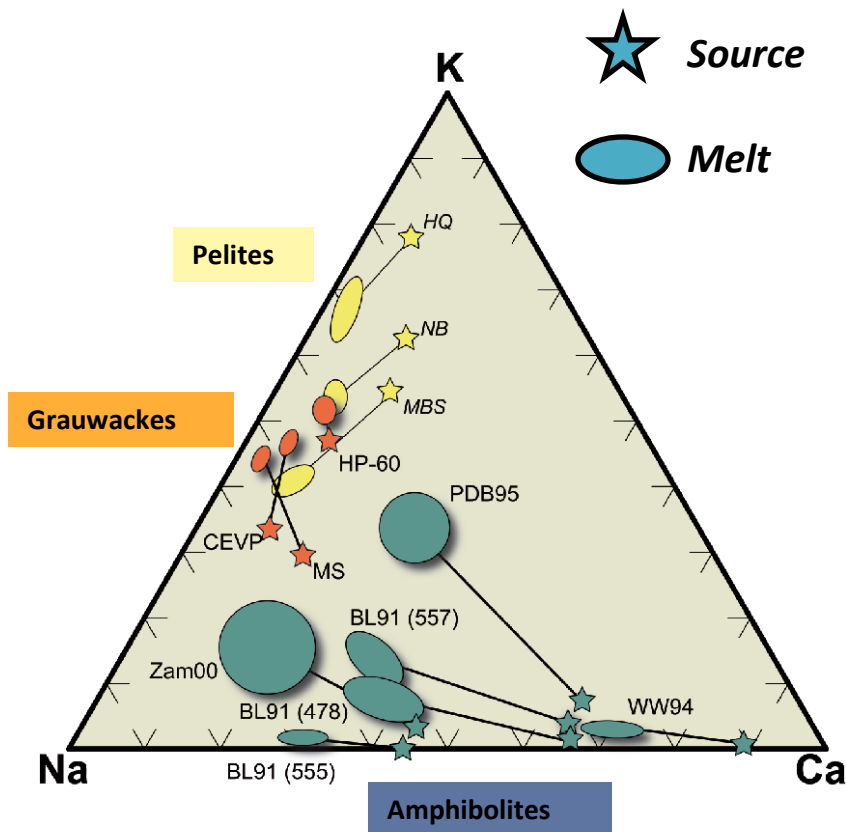
Pg + Qz + Bt + **KSp**

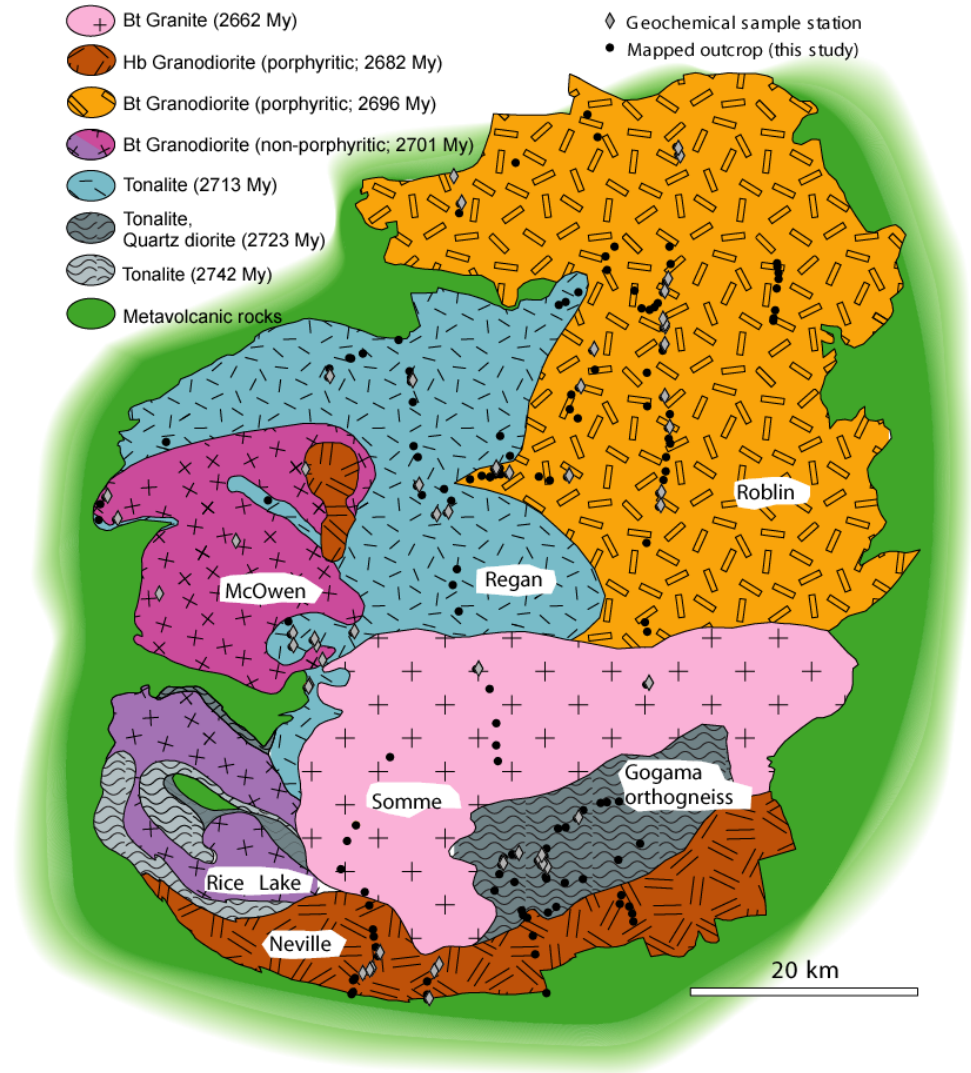
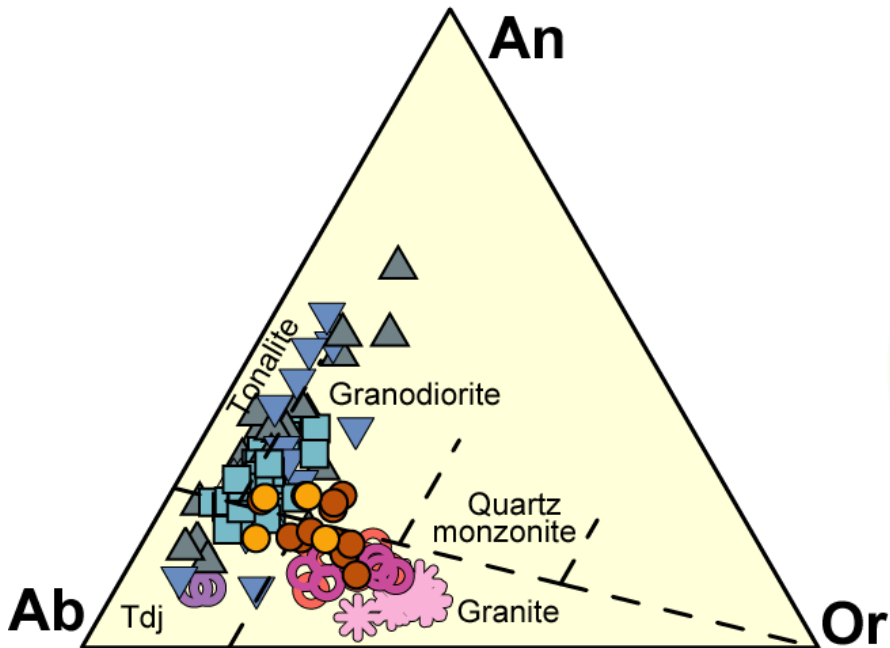
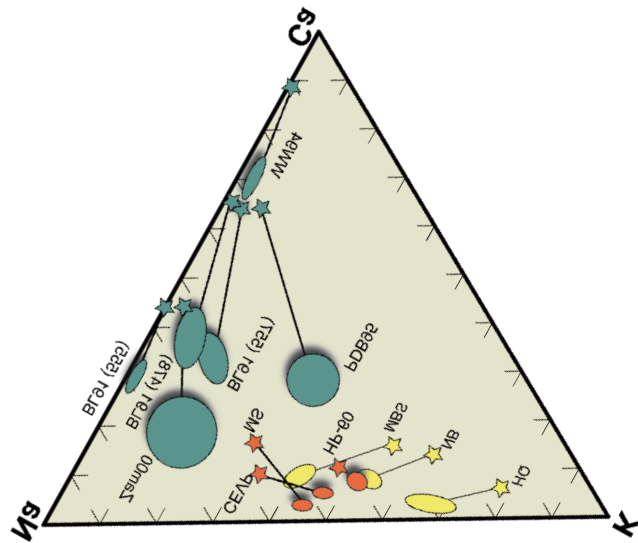
# C-types compositions



# Experimental melts and C-type granites

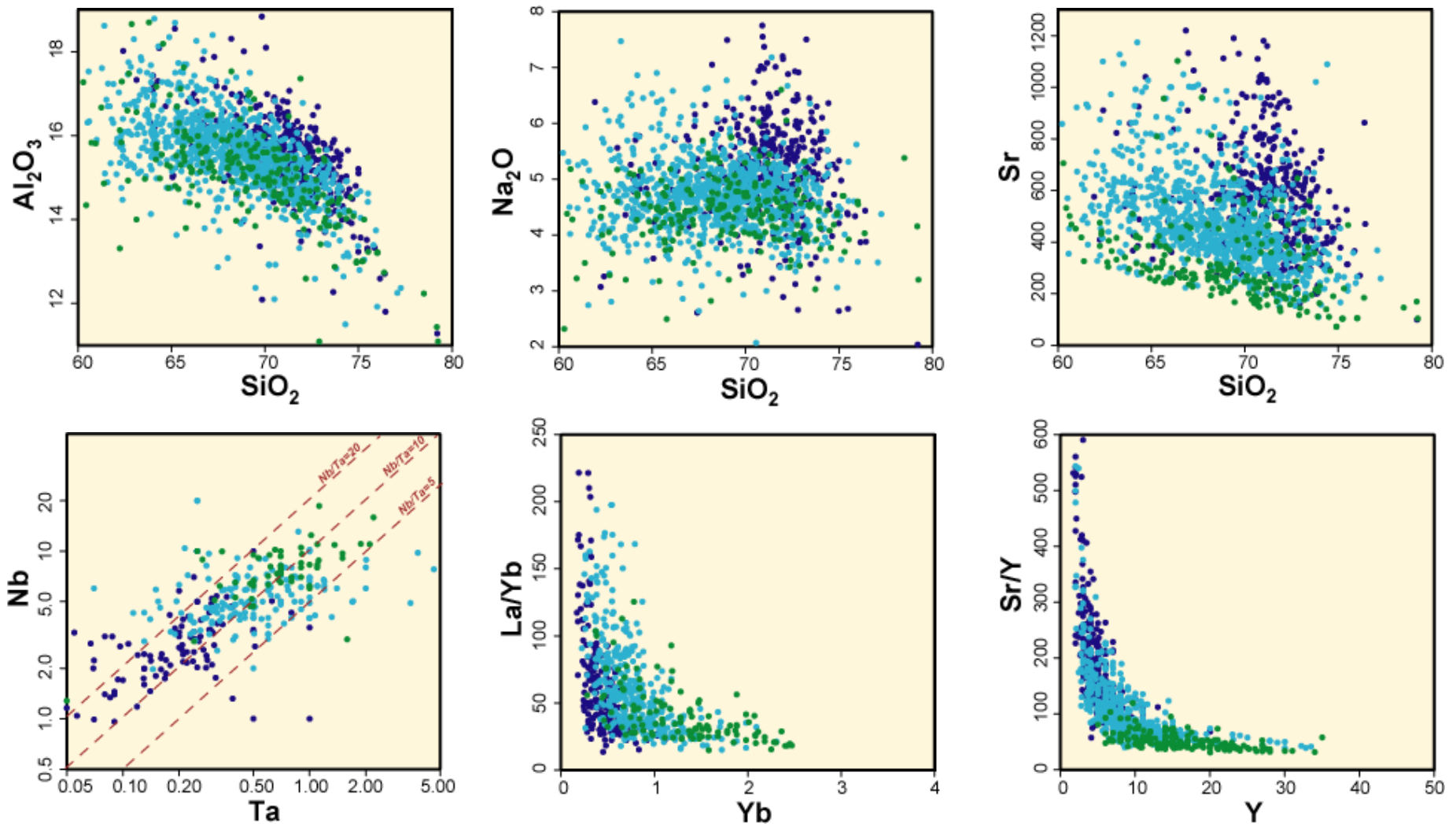
- Effect of different sources
- Effect of pressure (same source: Zamora, 2000)



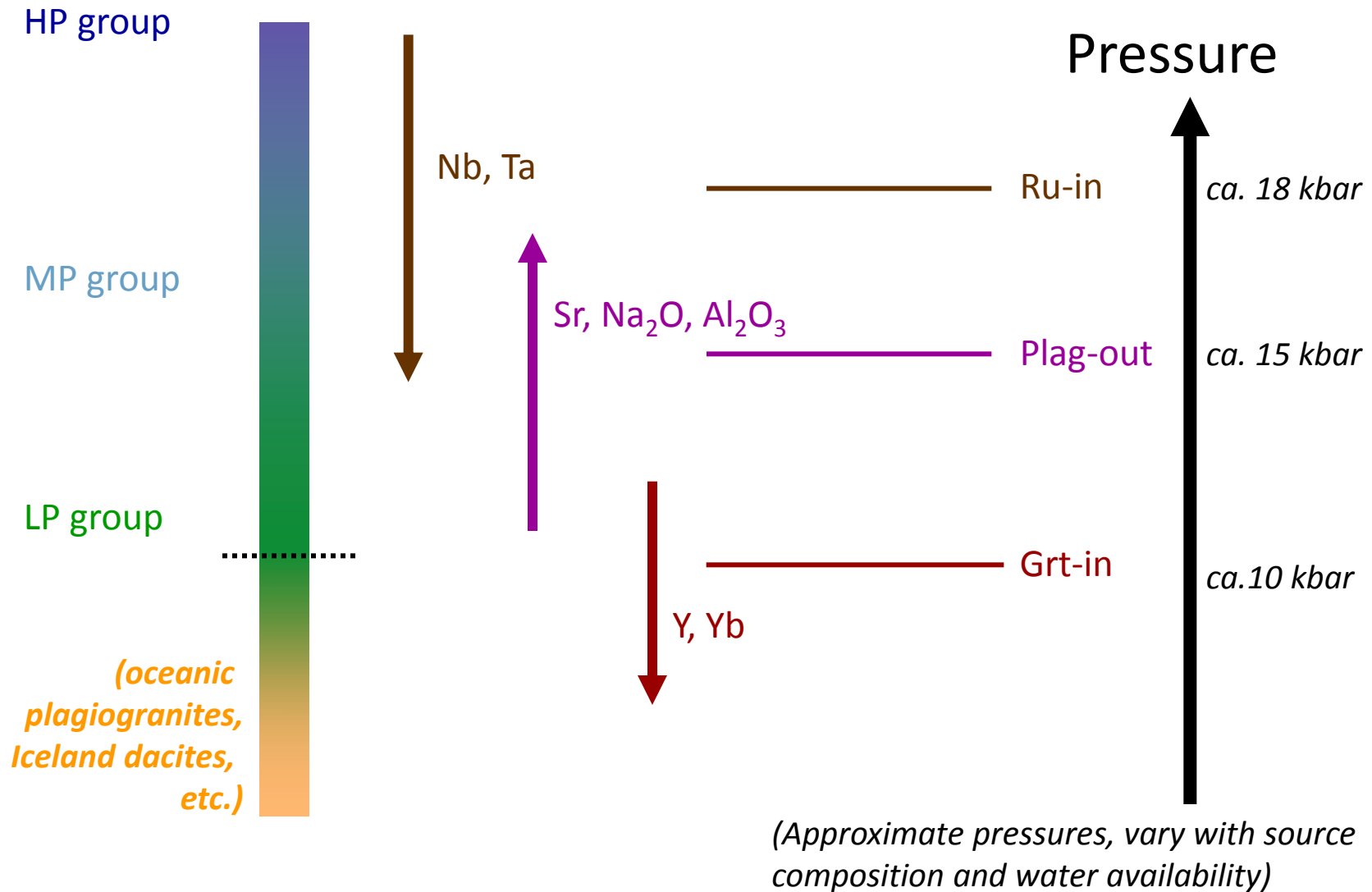


- + Bt Granite (2662 My)
- Hb Granodiorite (porphyritic; 2682 My)
- Bt Granodiorite (porphyritic; 2696 My)
- Bt Granodiorite (non-porphyritic; 2701 My)
- Tonalite (2713 My)
- Tonalite, Quartz diorite (2723 My)
- Tonalite (2742 My)
- Metavolcanic rocks
- ◆ Geochemical sample station
- Mapped outcrop (this study)

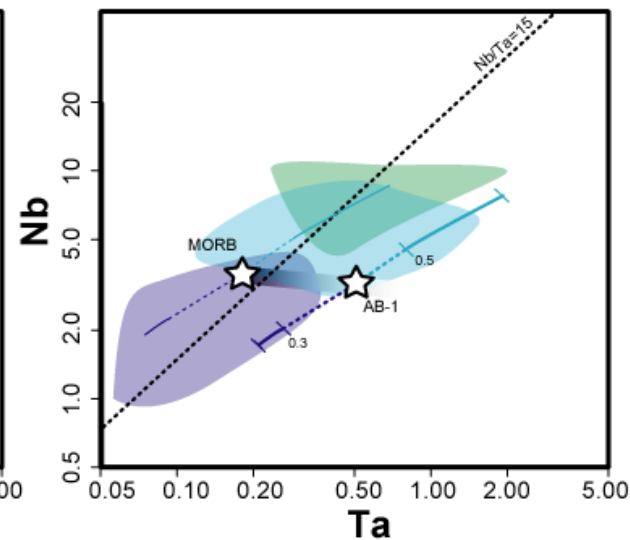
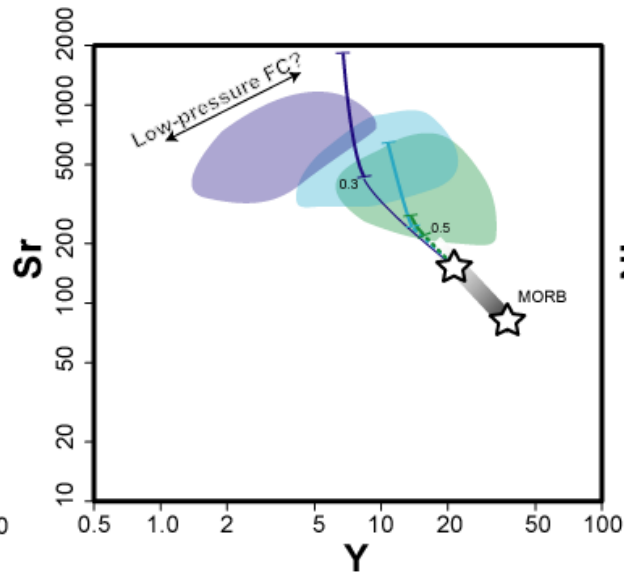
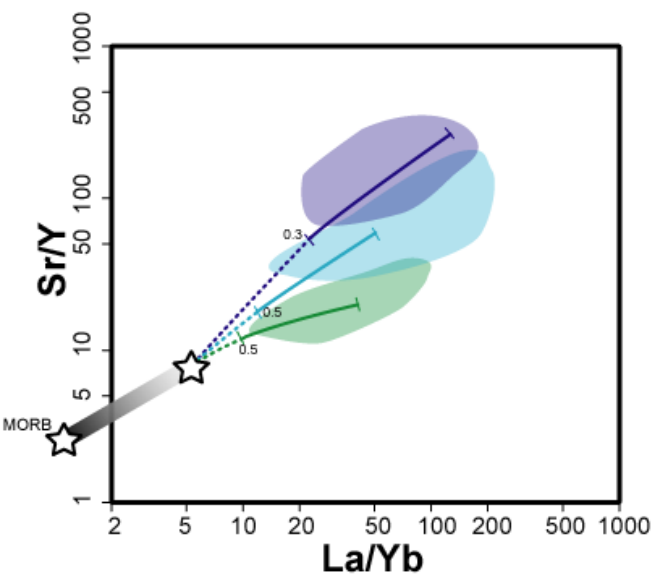
# TTG (sub)types – worldwide



# TTG subtypes: different pressures of melting



# Evidence for different melting depths in the TTG record



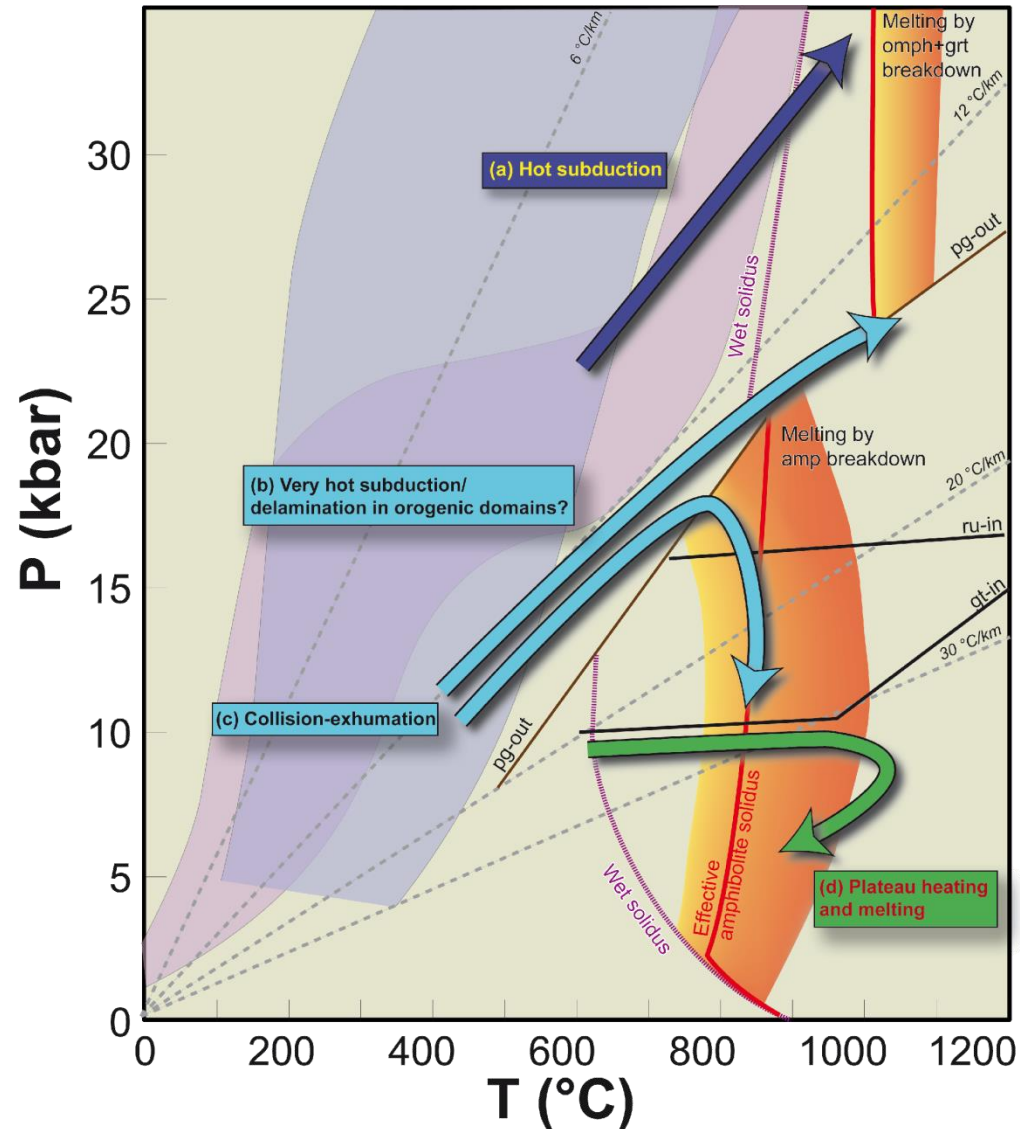
- |           |        |       |             |                  |
|-----------|--------|-------|-------------|------------------|
| 1. 20 Cpx | 50 Amp | 20 Pg | 10 Grt      | → ca. 10-12 kbar |
| 2. 25 Cpx | 60 Amp |       | 15 Grt      | → ca. 12-18 kbar |
| 3. 60 Cpx |        |       | 40 Grt 1 Rt | → ca. 20+ kbar   |

*Willbold, today 14.50*

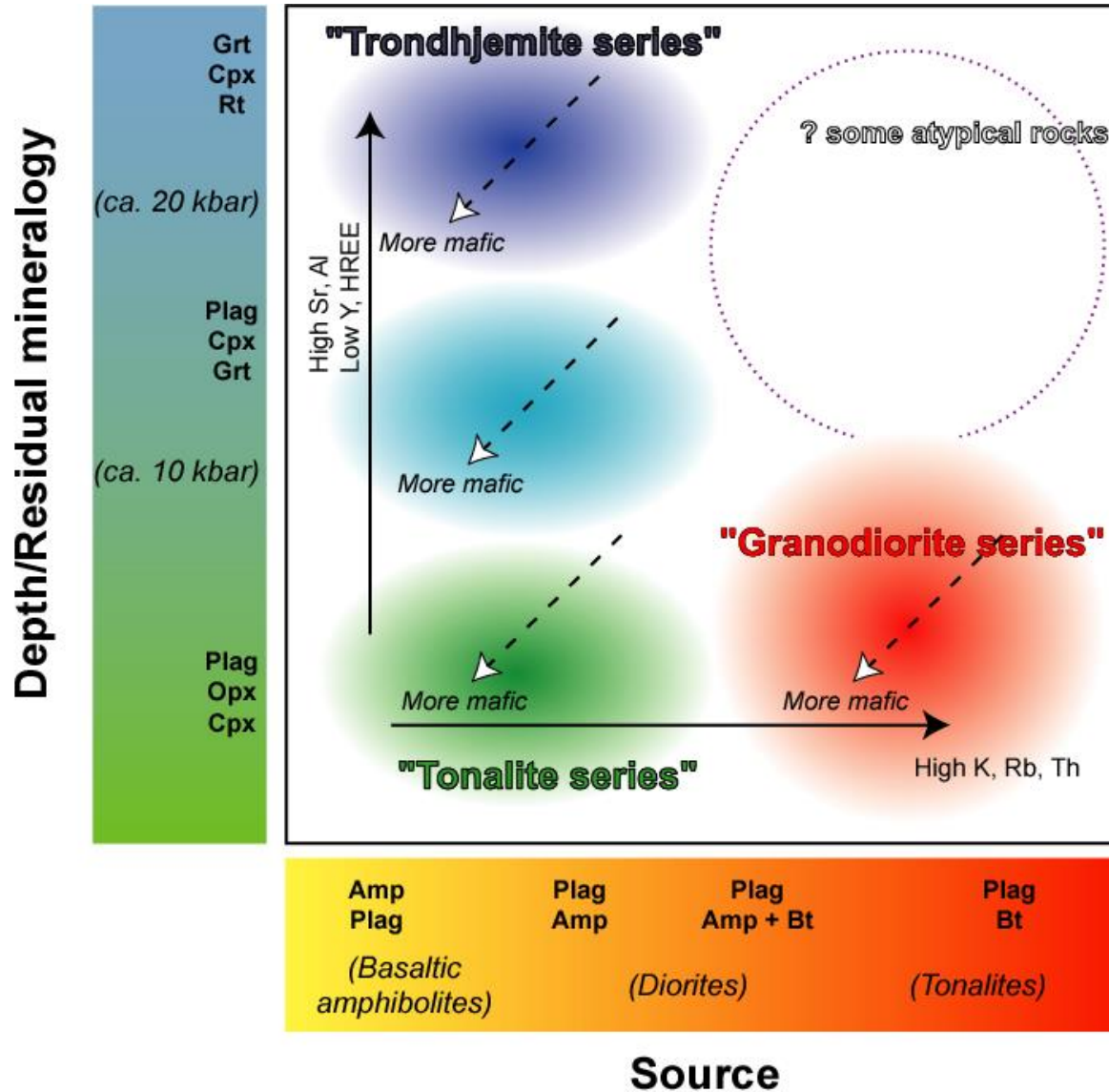
*Laurie, today 15.10*

*(Approximate pressures, vary with source composition and water availability)*

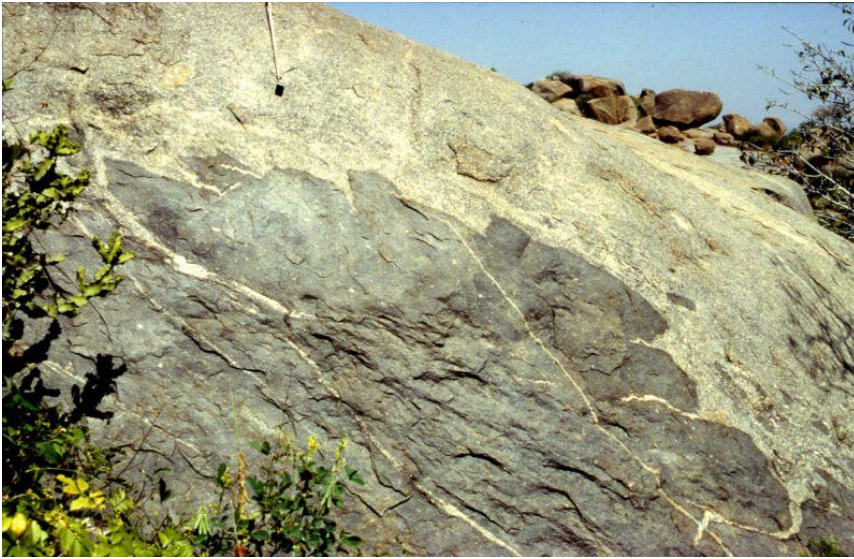
# PT conditions for different TTG types



# Source and melting pressure of Archaean C-type granitoids



# Archaean M-type granites (and associated intermediate diorites etc.)



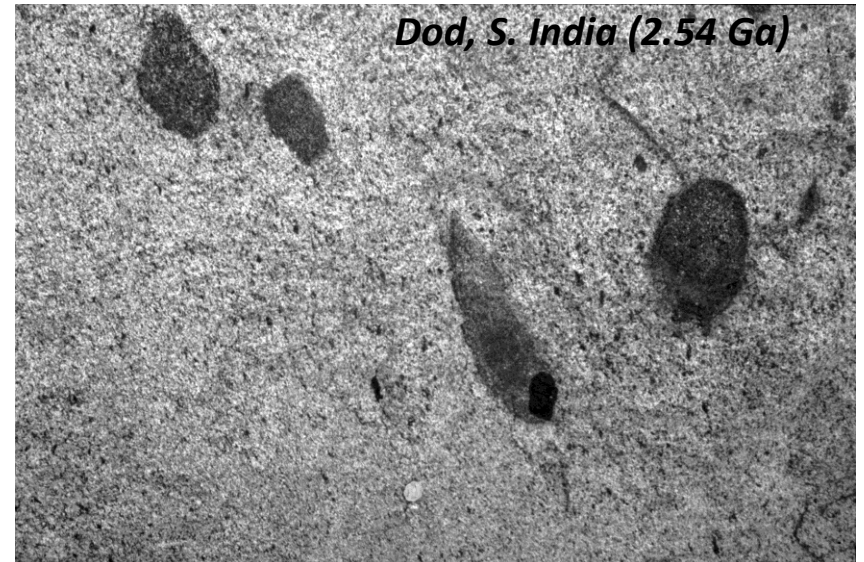
*Closepet, S. India (2.52 Ga)*



*Peewah, WA (2.95 Ga)*



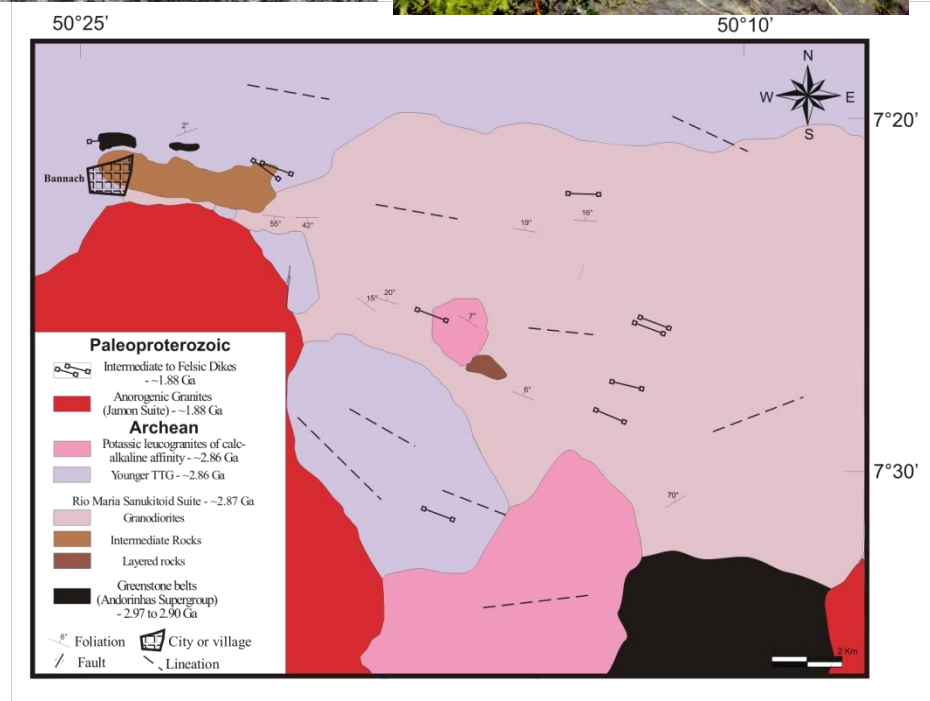
*Matok, S. Africa (2.67 Ga)*



*Dod, S. India (2.54 Ga)*

# M-type granitoids: link with intermediate rocks (K-diorites)

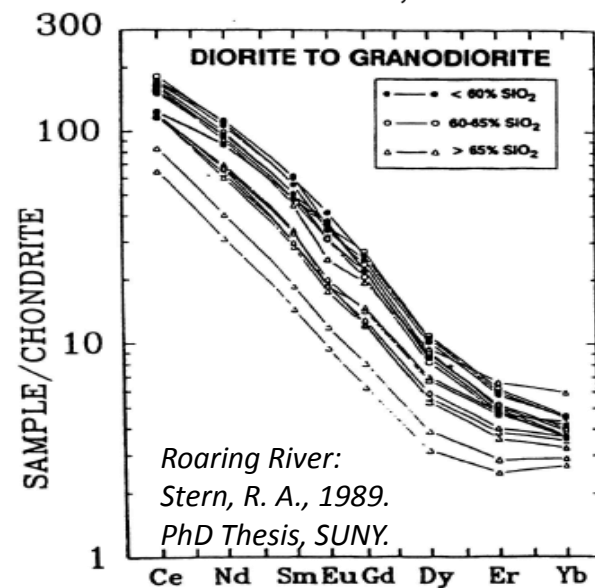
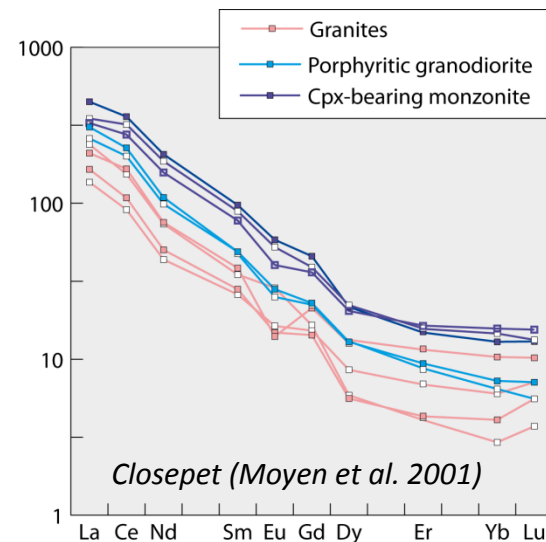
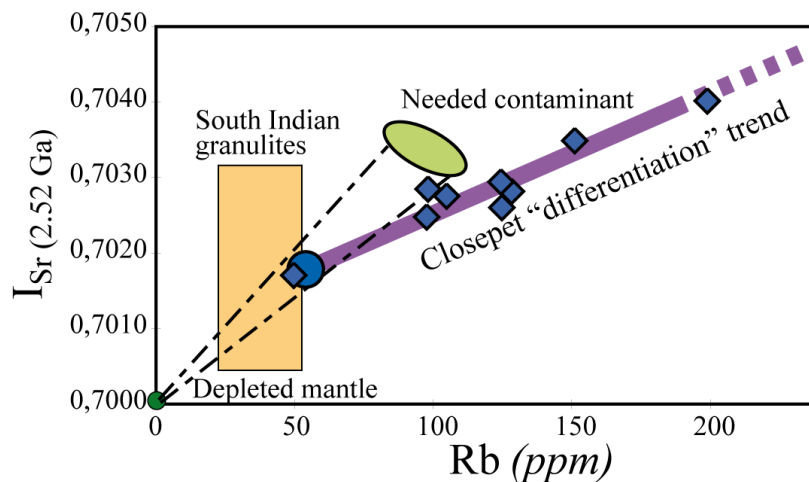
- M-types are always associated with intermediate rocks
  - Evidenced by MME
  - But also larger stocks up to  $\approx 1$  km
- K-diorites
- 2 petrological questions:
  - Relations between diorites and granodiorites/granites
    - Fractional crystallization?
    - Remelting of mafic precursor?
    - In-situ mixing with local country rocks or crustal melts?
  - Origin of the diorites



Rio Maria GGT, Oliveira et al.

# M-type granitoids: origin of the K-diorites

- The most enriched components in the suite are the diorites
- They typically have higher LILE & LREE than surrounding crust
  - Crustal contamination unlikely
  - Enriched mantle source



## sanukitoid origins

Melting of TTG-metasomatized mantle III:  
dehydration of amphibole, w/out garnet signature  
(Amphibole + OPX  $\pm$  CPX)

P (GPa) 1.6  
Temp (°C) 1100  
Melt/Rock 1:2

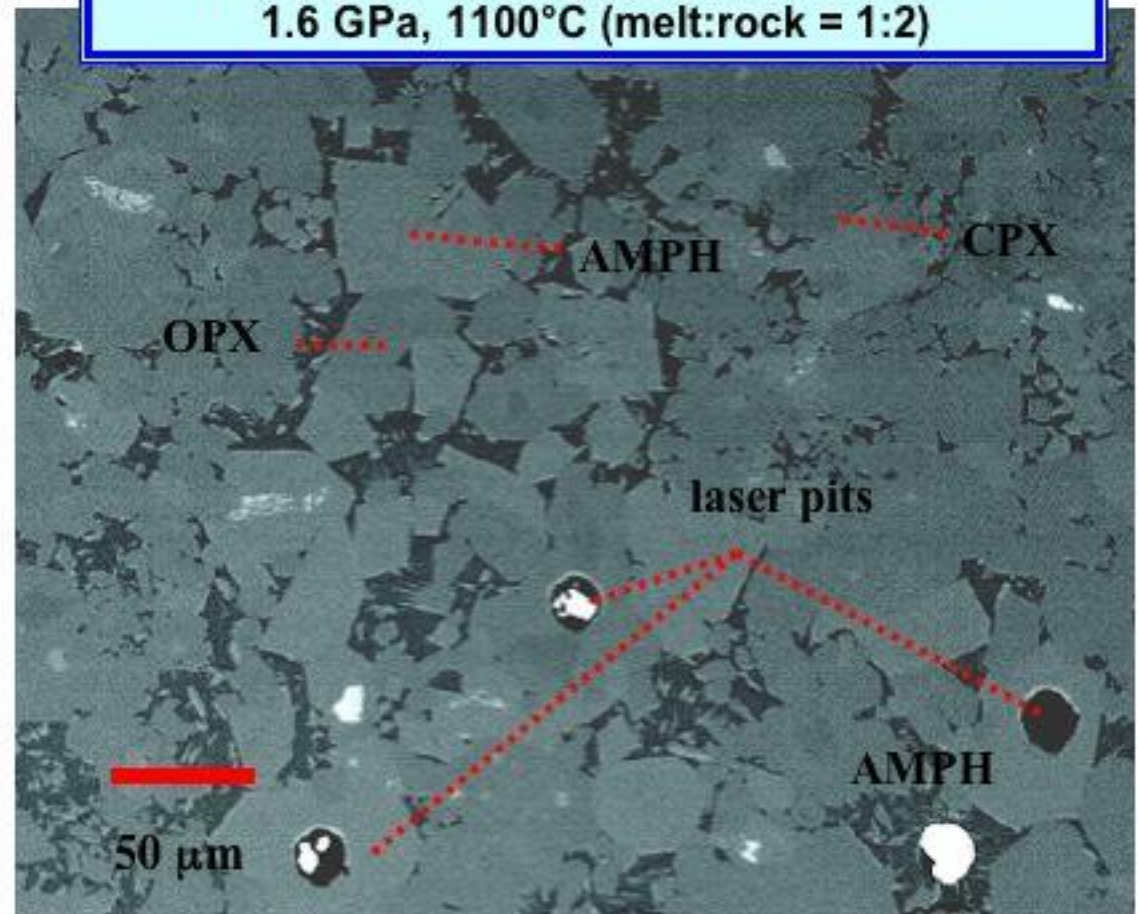
SiO<sub>2</sub> 54.35  
TiO<sub>2</sub> 0.38  
Al<sub>2</sub>O<sub>3</sub> 18.60  
FeO 3.43  
MnO 0.05  
MgO 2.59  
CaO 3.63  
Na<sub>2</sub>O 6.05  
K<sub>2</sub>O 2.23

TOTAL 91.64%  
Mg# 0.57

P = 1.6 GPa

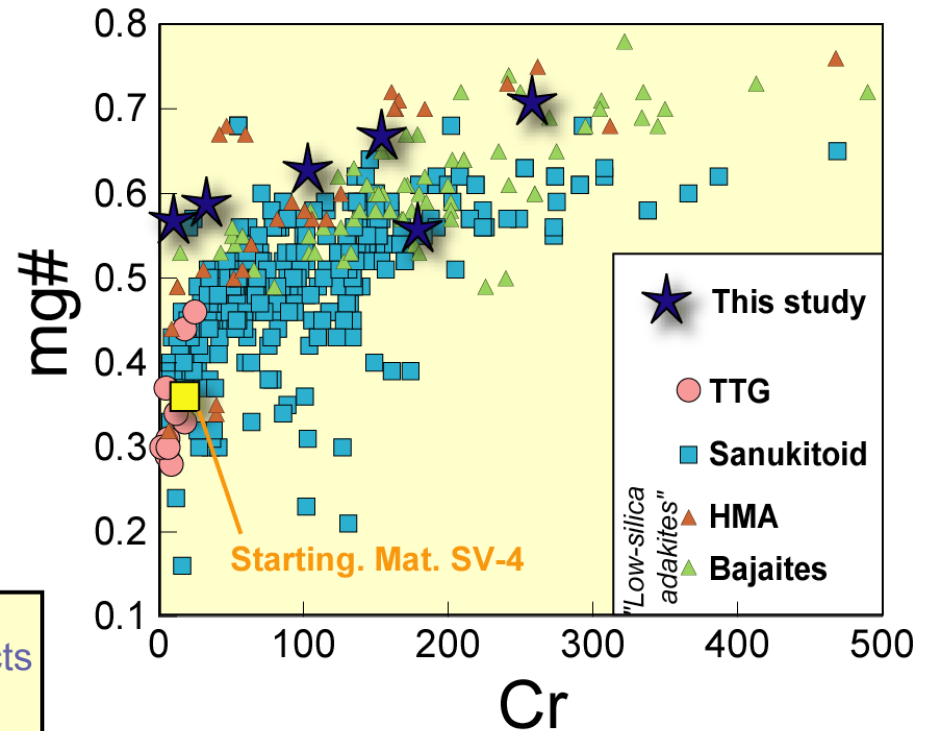
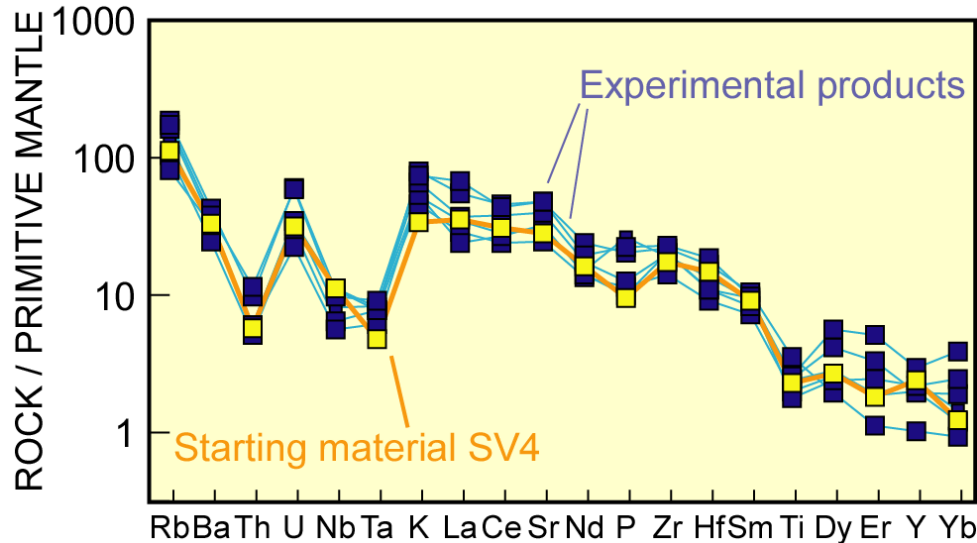
*amphibole signature,  
OPX signature*

Hydrous adakite SV-4 + peridotite 8710P  
1.6 GPa, 1100°C (melt:rock = 1:2)

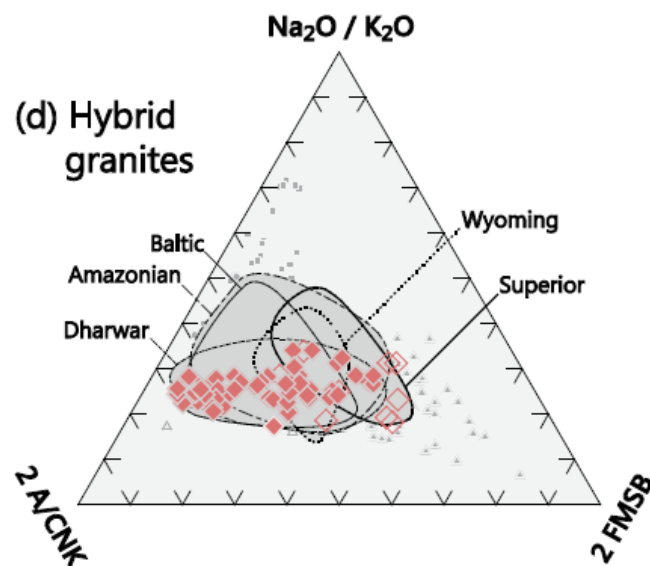
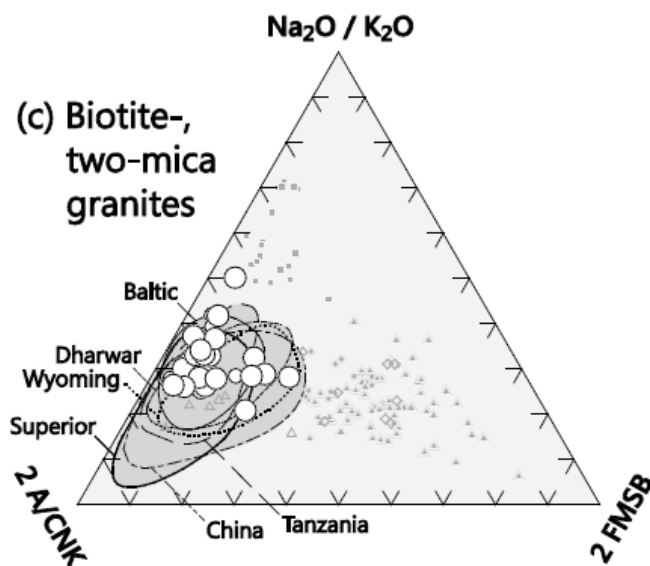
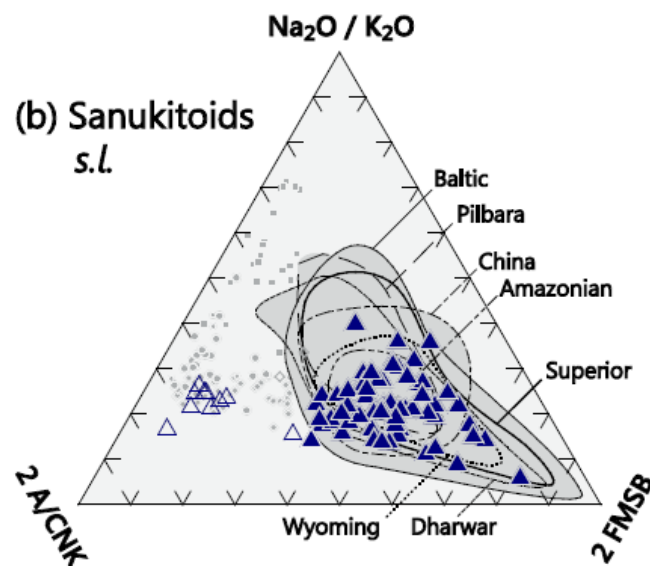
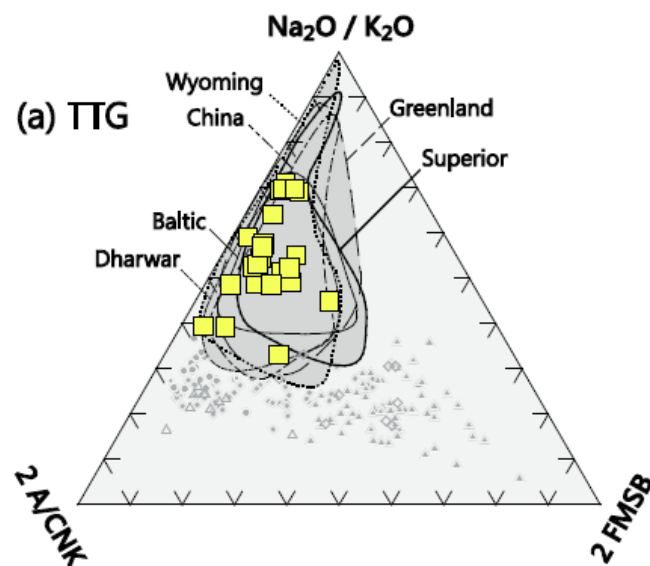


# Experimental results

As expected, experimental melts have high mg# and Cr, and TTG-like trace element patterns are preserved.

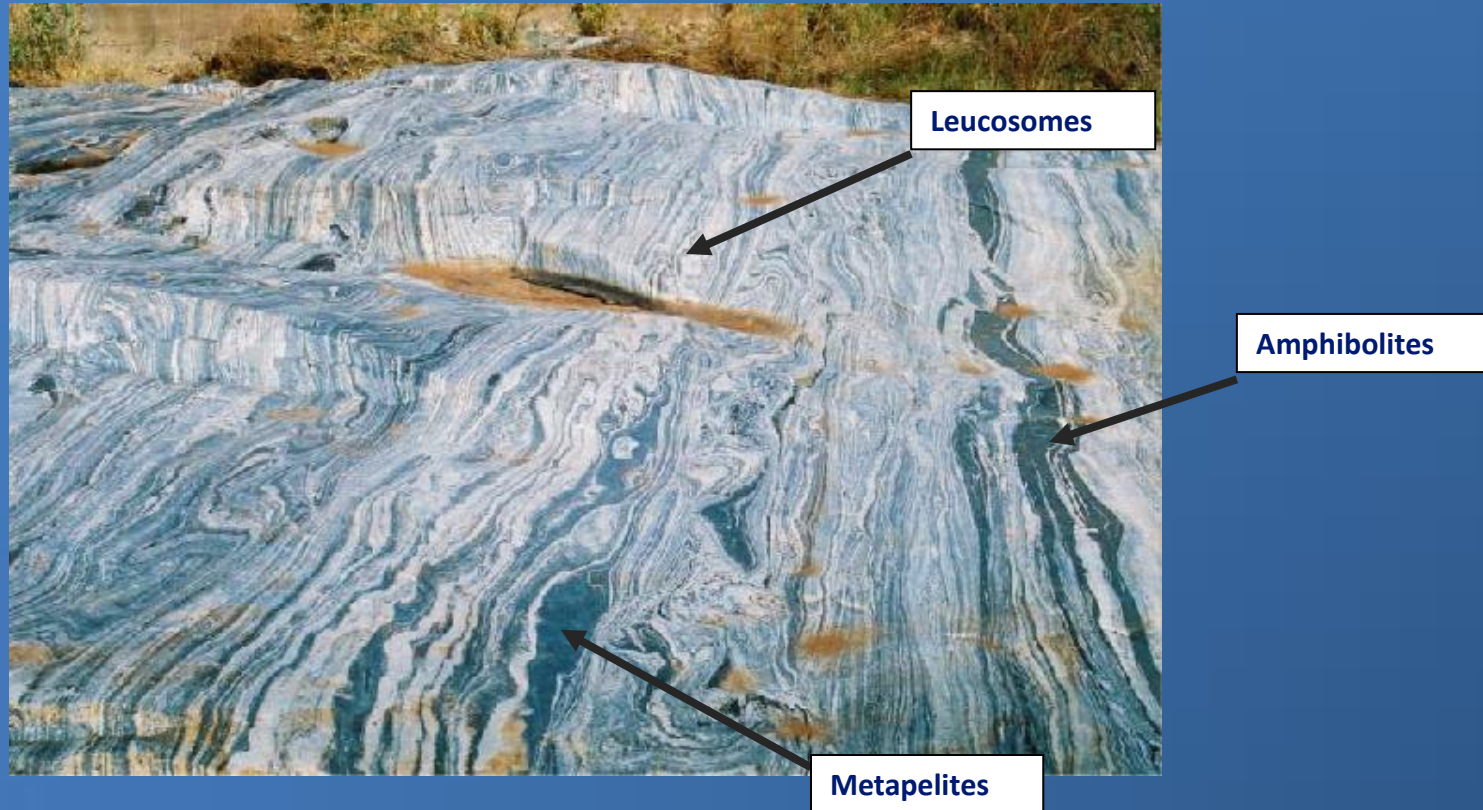


# Hybrid granites



# Granites, gneiss et migmatites

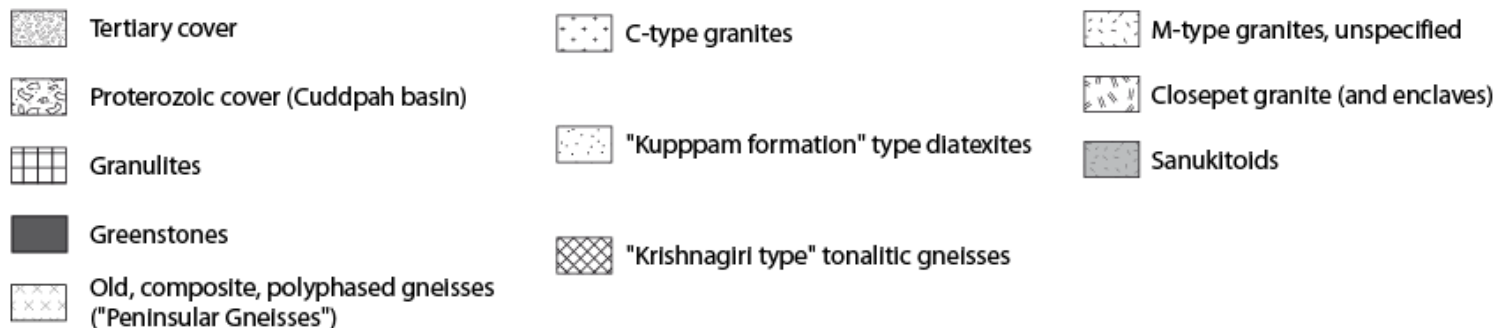
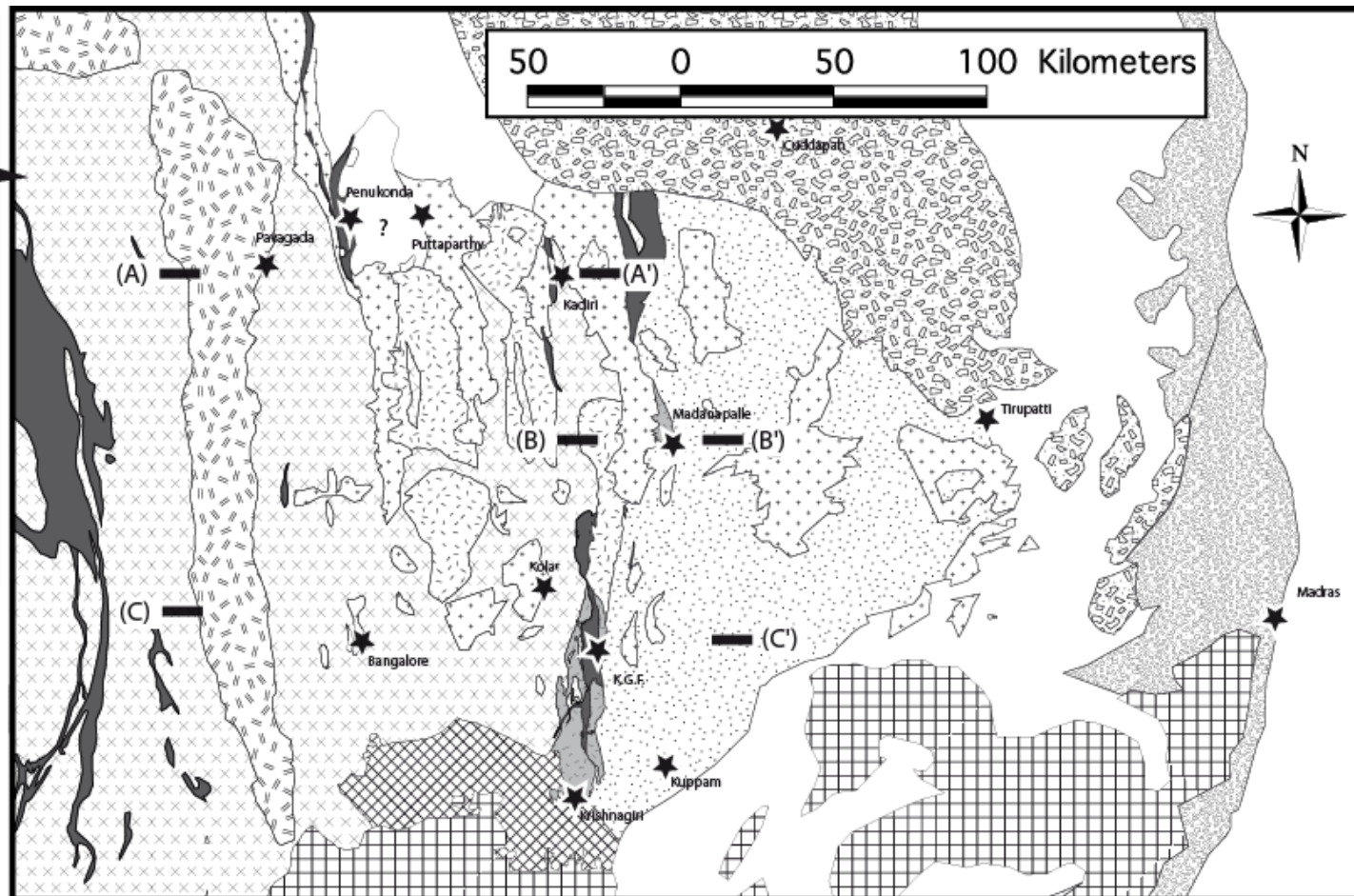
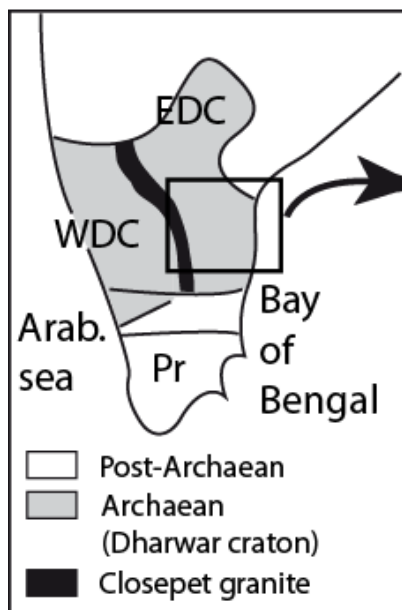
*Sand River Gneisses (3.1 Ga), Limpopo, South Africa*

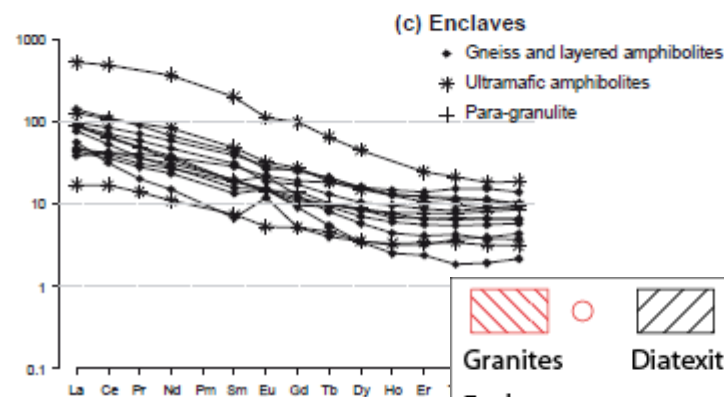
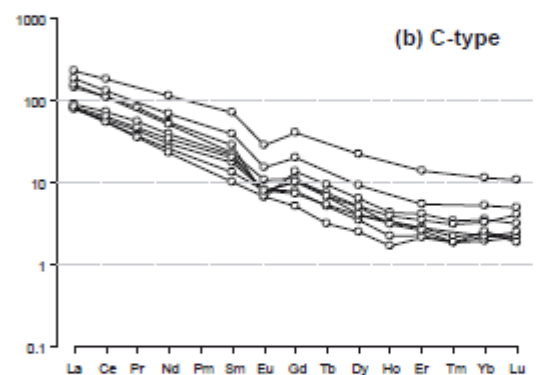
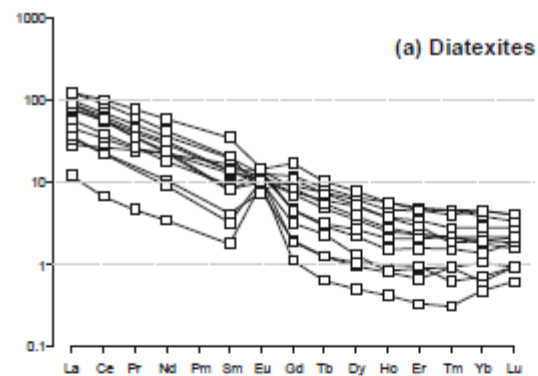


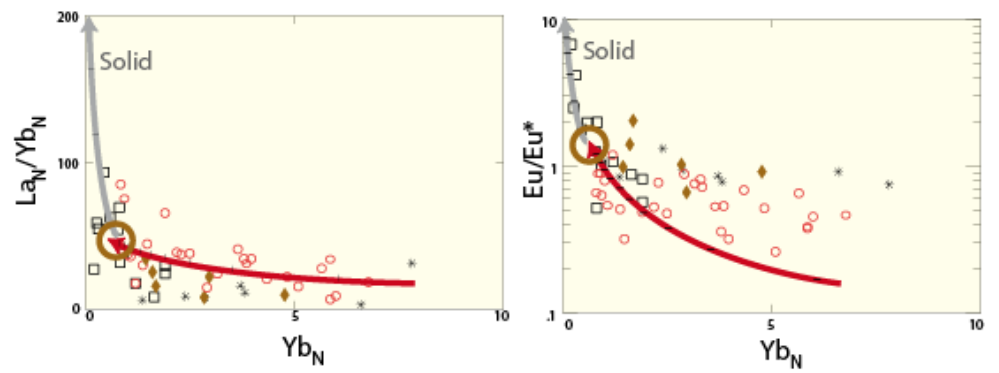
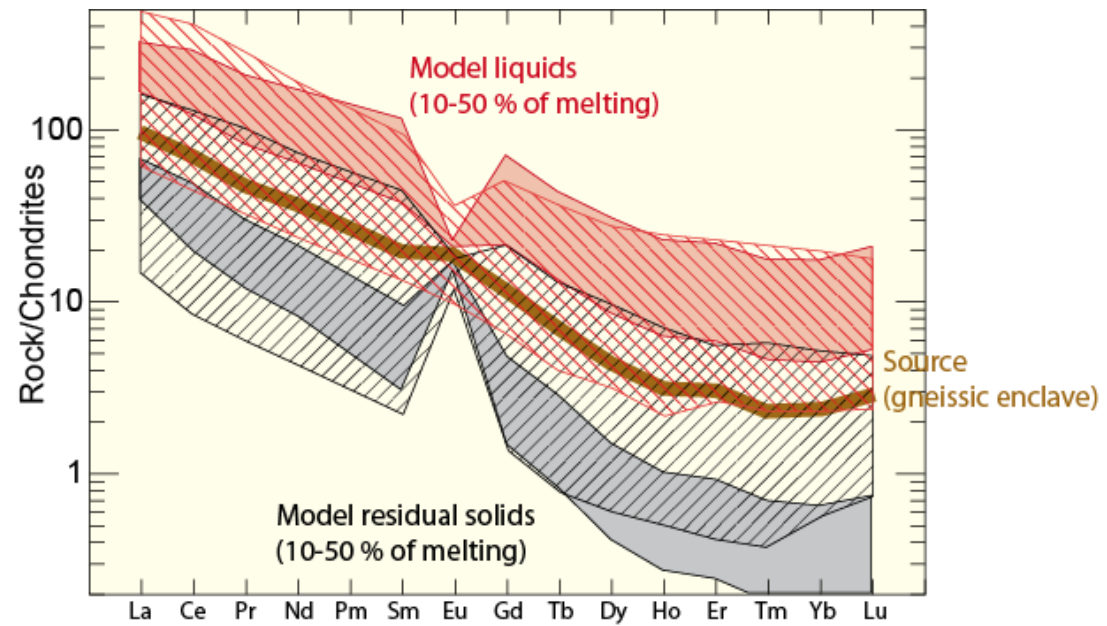
- **Grey gneisses** = « background » lithology of Archaean terranes. May form up to 80% of Archaean C.C.
- **TTG** = Tonalite-Trondhjemite-Granodiorite. **Sodic** granitoids, high Sr/Y and La/Yb, form by melting of metamafics. Supposedly protolith of grey gneisses (!)

# Grey Gneisses and TTG

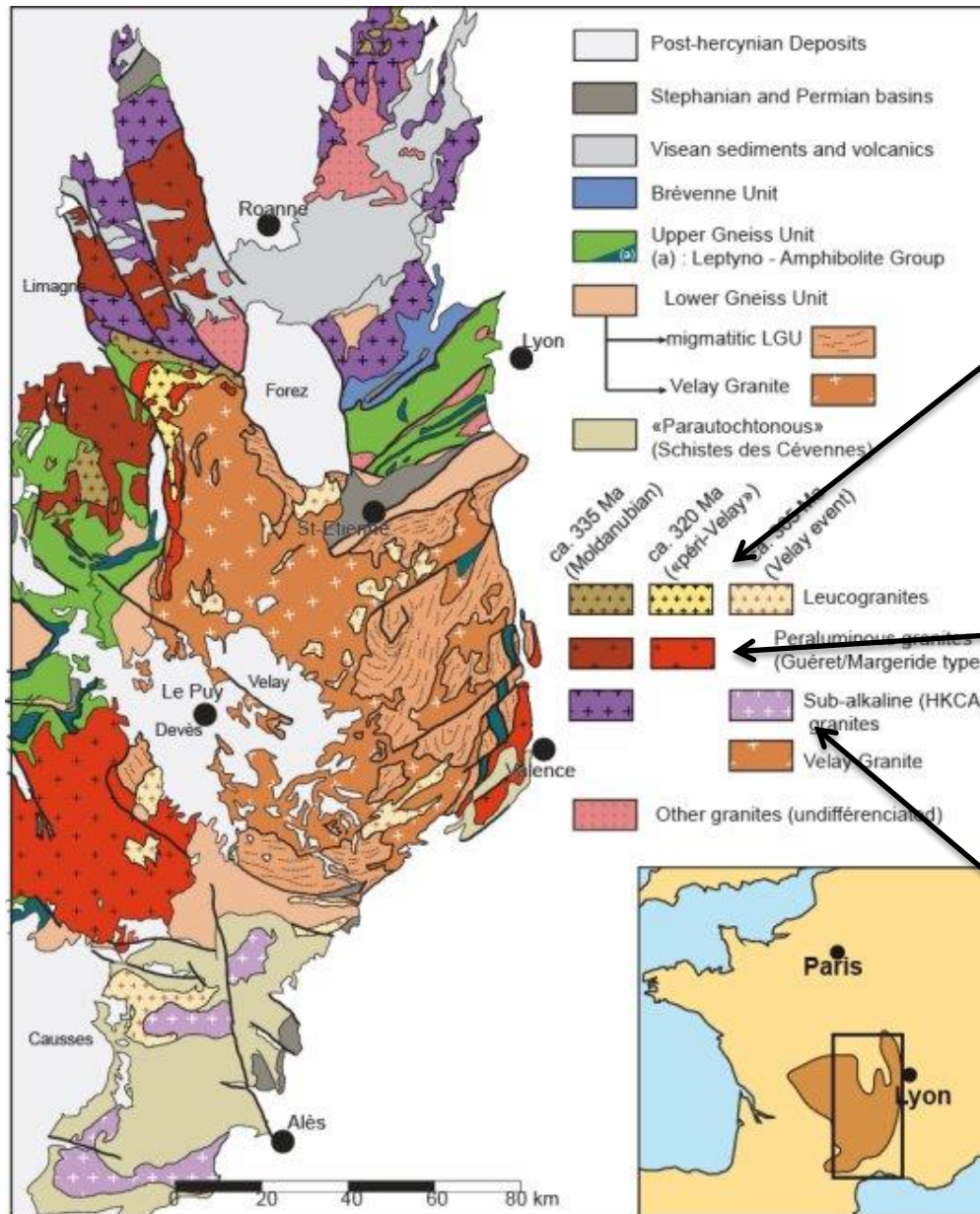




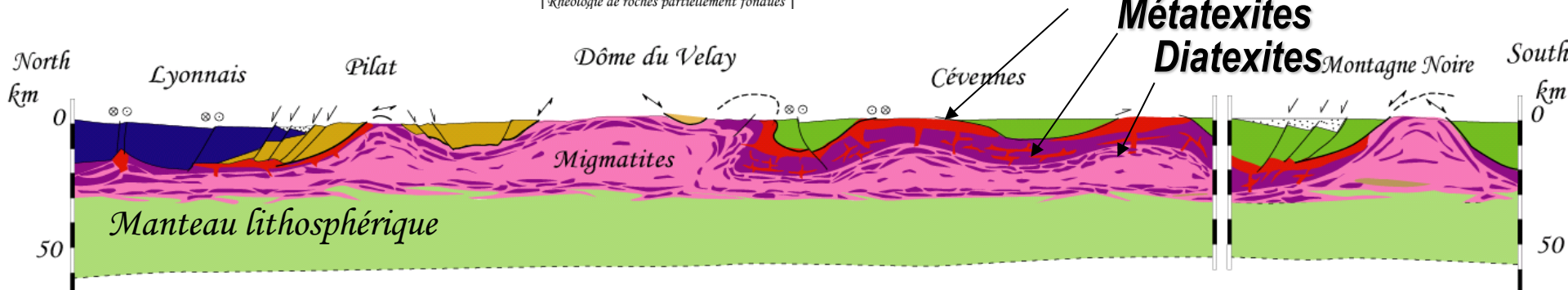
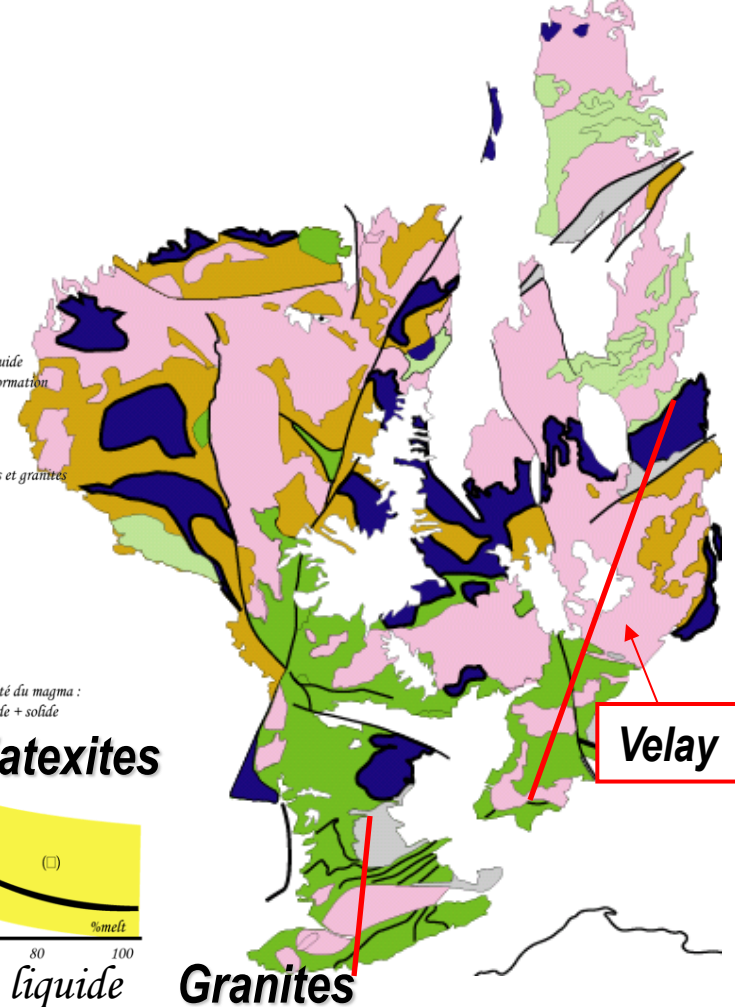
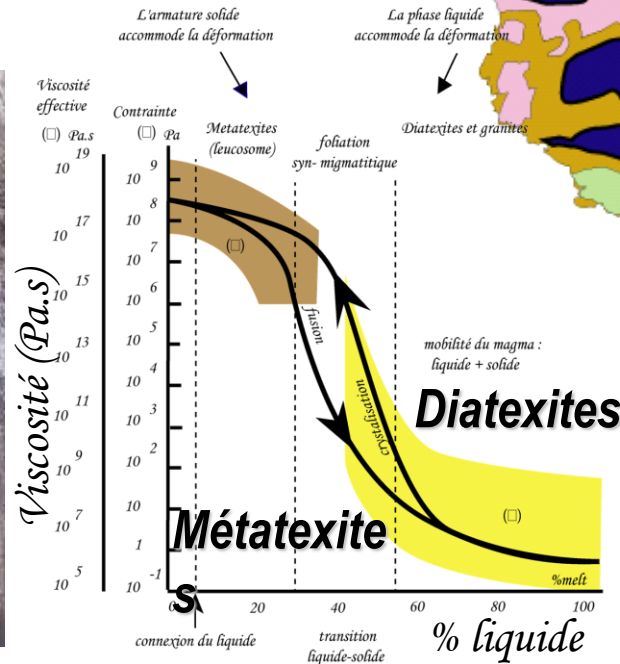
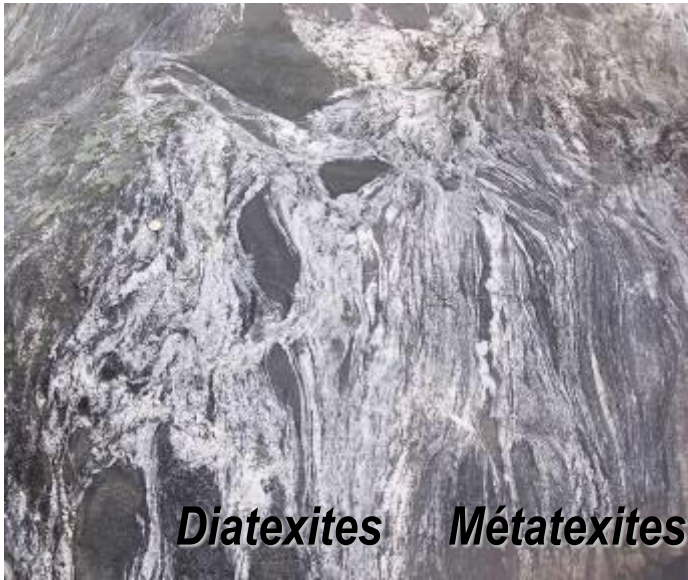




# Le Massif Central Français: granites

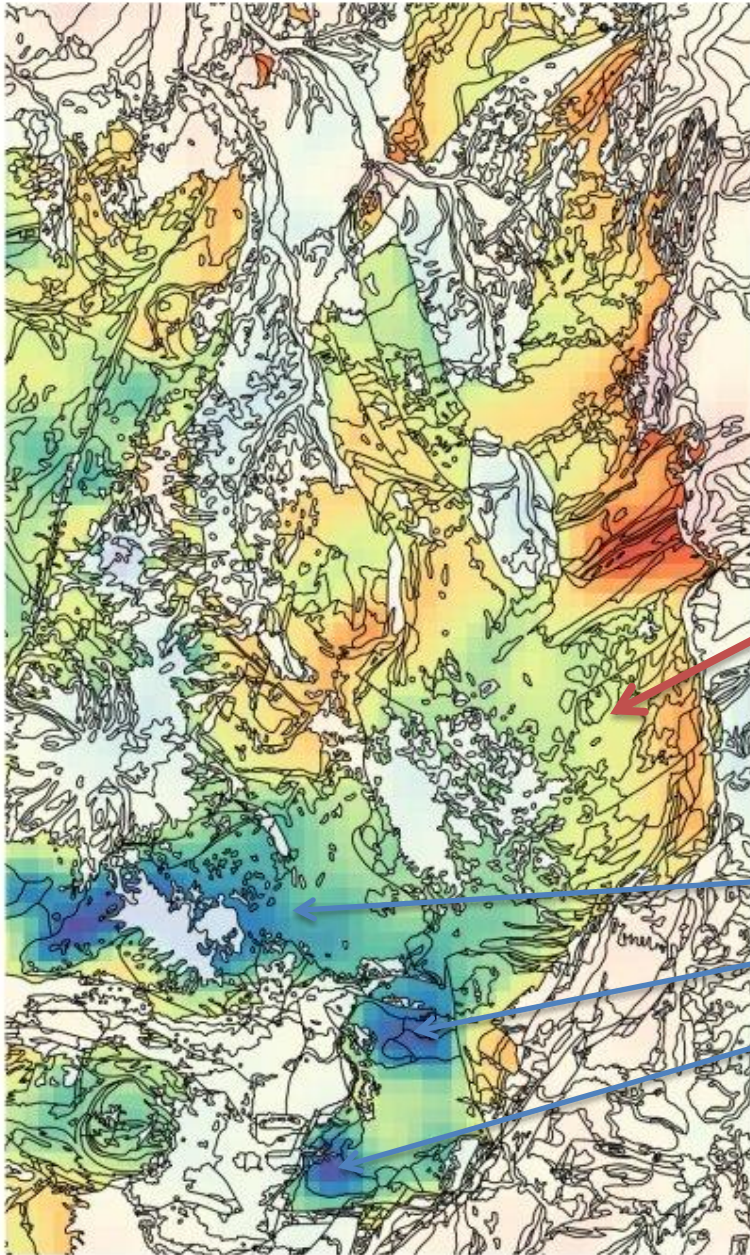


# ..et migmatites



25 0 25 50 75 100 km

# Bouguer anomaly

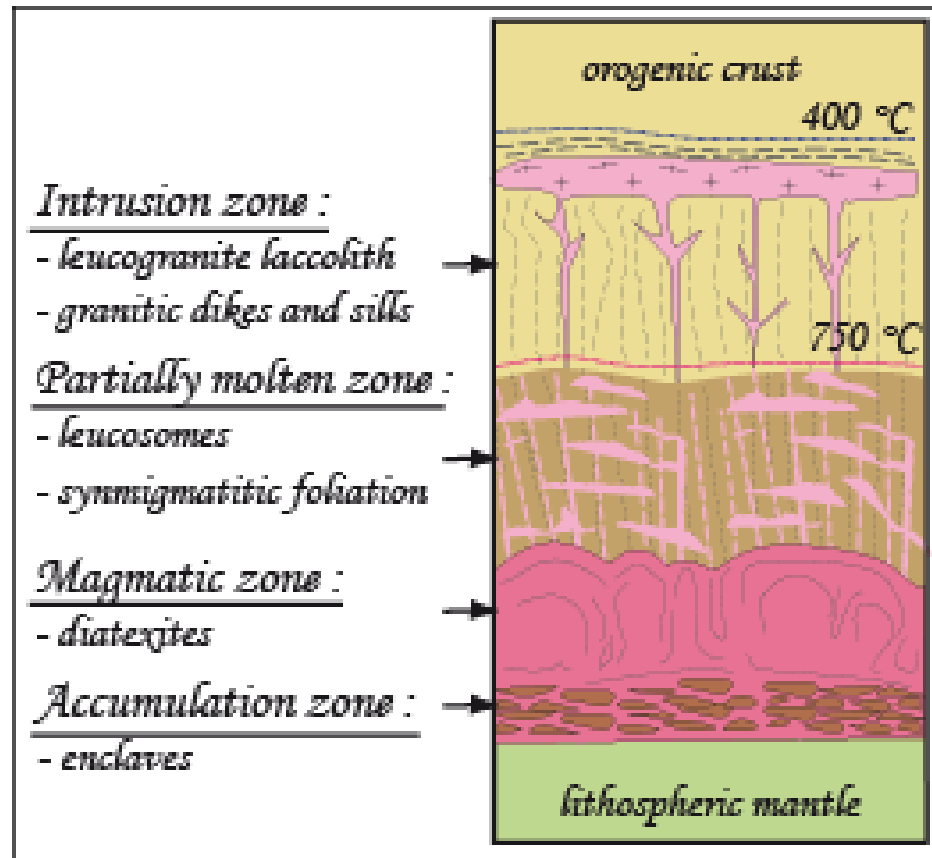


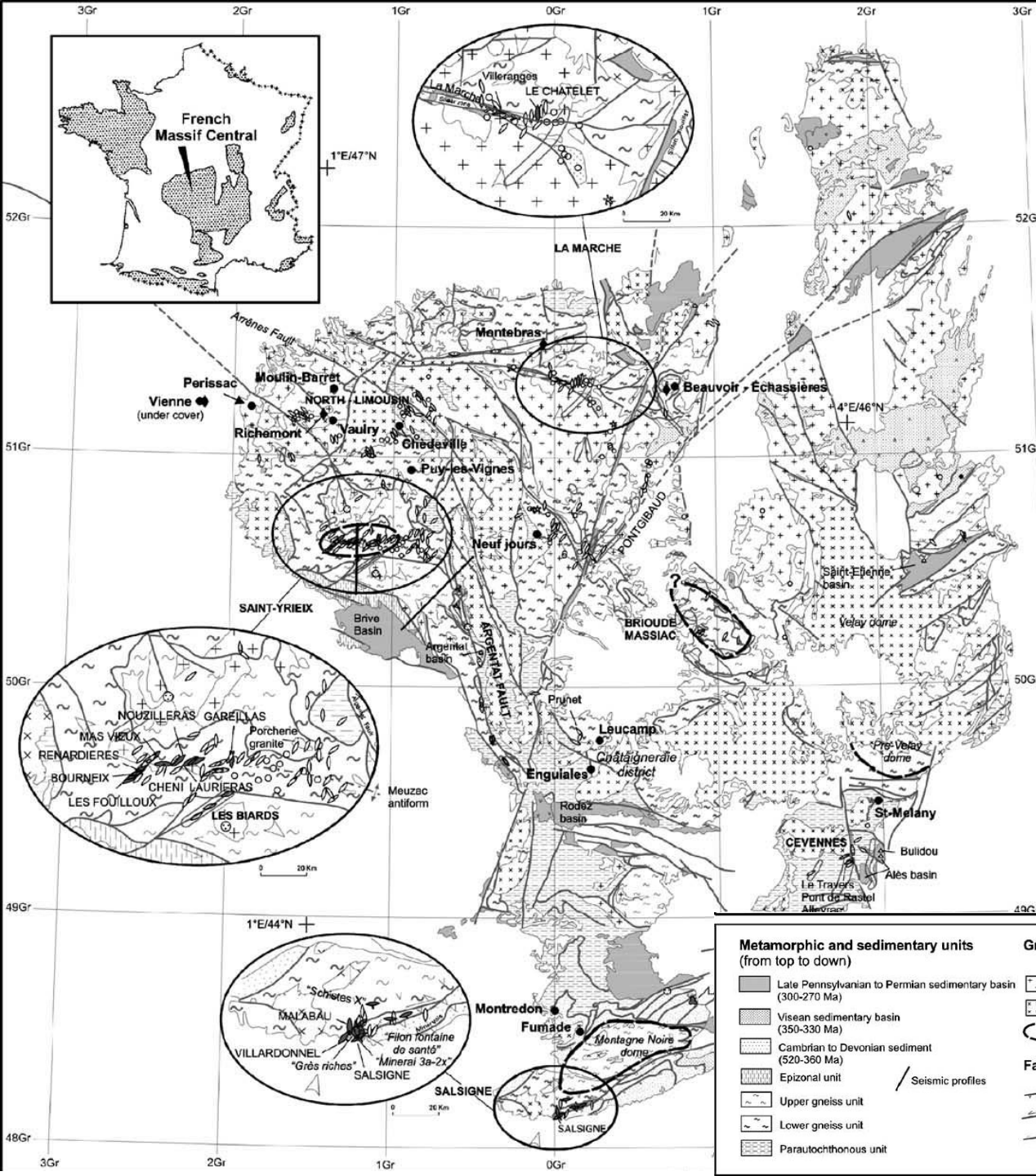
Velay complex = no melt extracted  
Source layer outcrops  
***No Bouguer anomaly***

Granites = melt extracted  
from source layer  
***Negative Bouguer anomaly***



# Structure de la croûte anatectique





#### Metamorphic and sedimentary units (from top to down)

- Late Pennsylvanian to Permian sedimentary basin (300-270 Ma)
- Visean sedimentary basin (350-330 Ma)
- Cambrian to Devonian sediment (520-360 Ma)
- Epizonal unit
- Upper gneiss unit
- Lower gneiss unit
- Parautochthonous unit

#### Granitoids

- > 335 Ma granitoid
- < 335 Ma granite
- 335-315 Ma granito-migmatite dome

#### Faults

- Thrust
- Strike-slip fault
- Normal fault

#### Expression of gold mineralization

- Vein type
- Stockwork type
- Sulfide dissemination type
- Skarn type
- Paleoconglomerate type

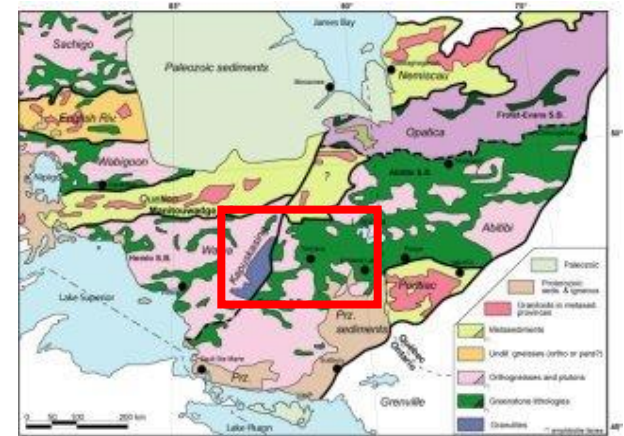
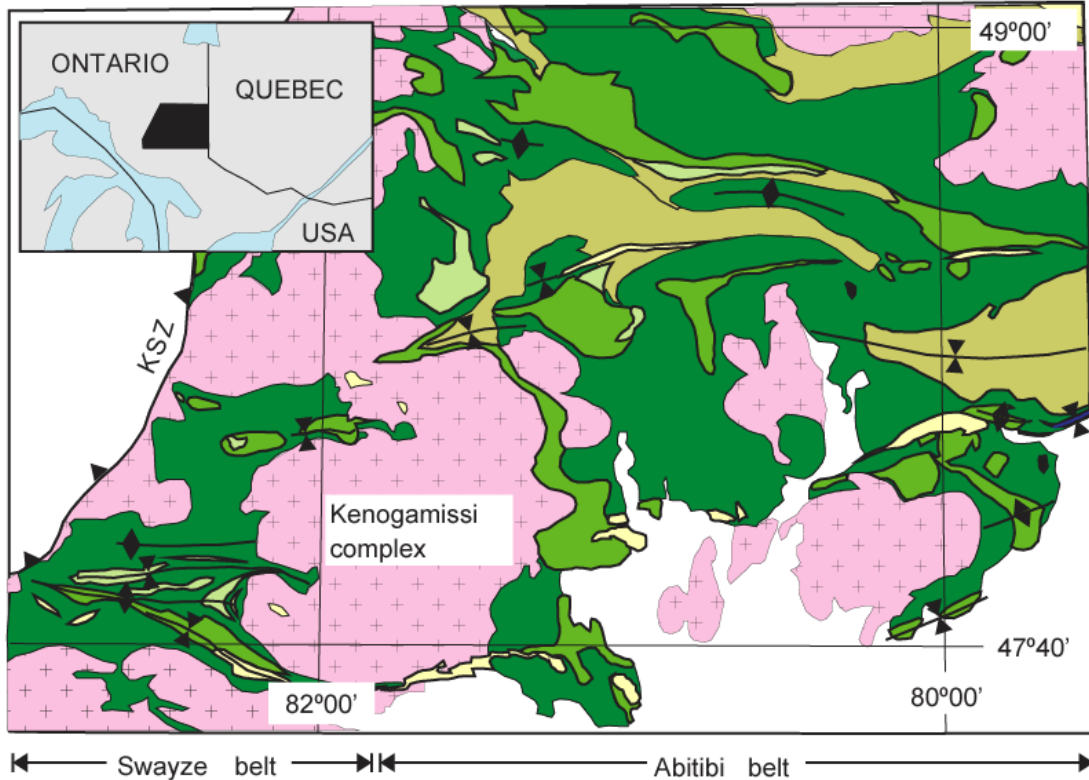
#### Size of gold mineralization

- Deposit
- Occurrence
- SAINT-YRIEX** = Au district name
- BOURNEIX** = Au deposit name

#### Granite related mineralization

- Beauvoir** { W + Sn deposit
- Rare-metal deposit

# Abitibi: zone des plutons?



Granite  
2740-2660 Ma

## Metasedimentary rocks

Wacke, Conglomerate 2700-2685 Ma

Wacke, Sandstone ca. 2700 Ma

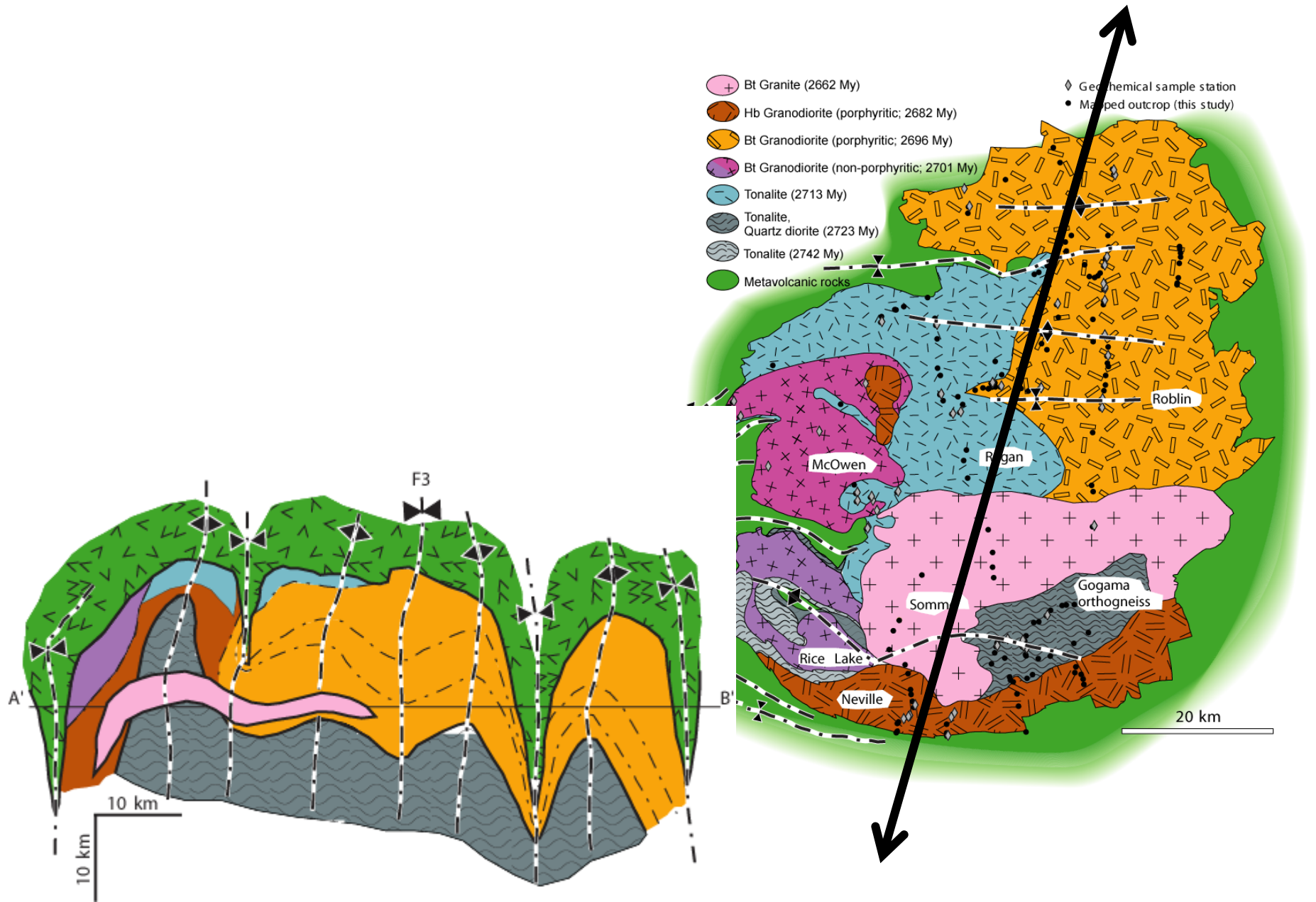
## Metavolcanic rocks 2750-2700 Ma

Felsic

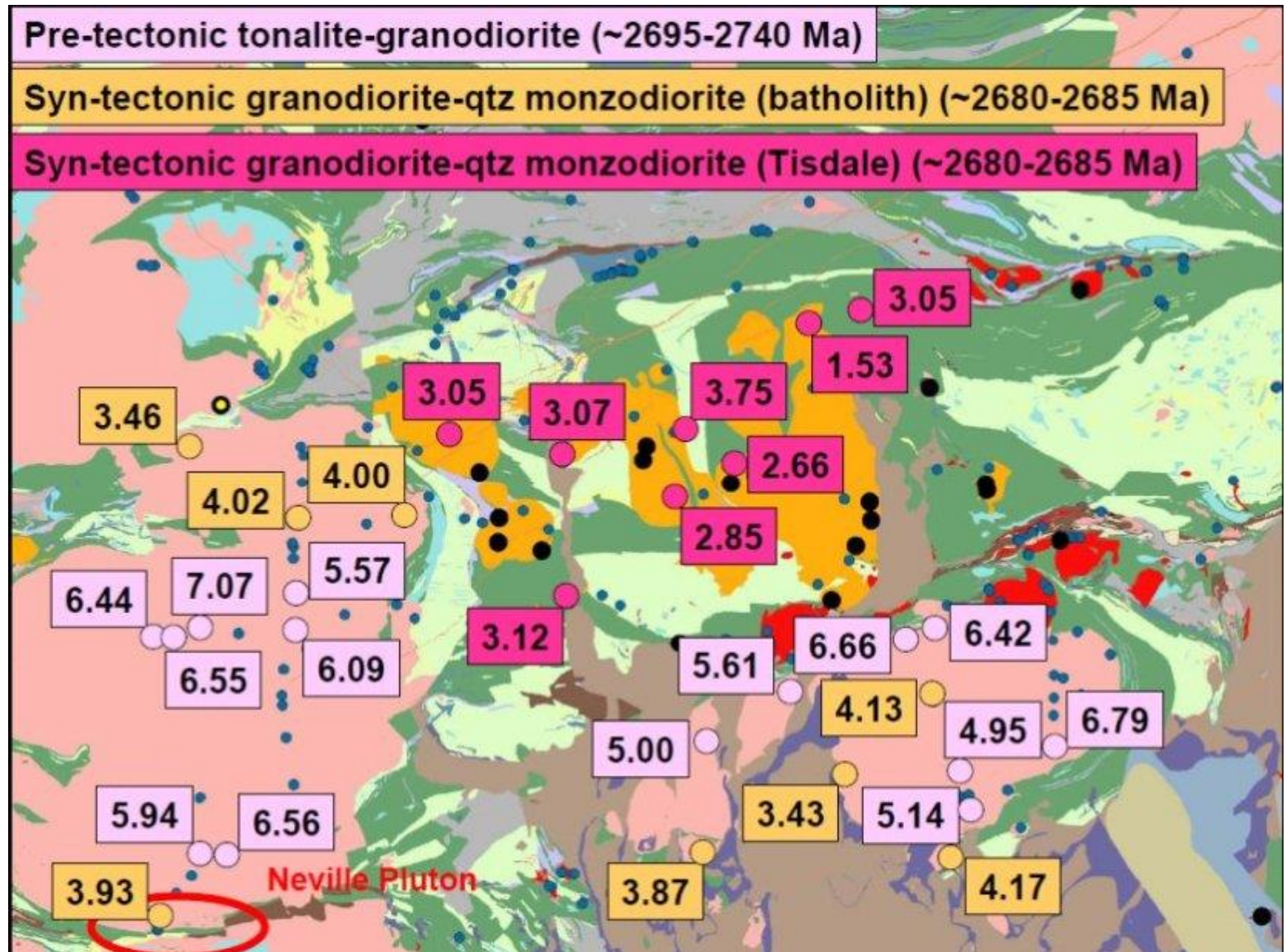
Intermediate

Mafic

# Kenogamissi complex: structure



# Map patterns and crustal level

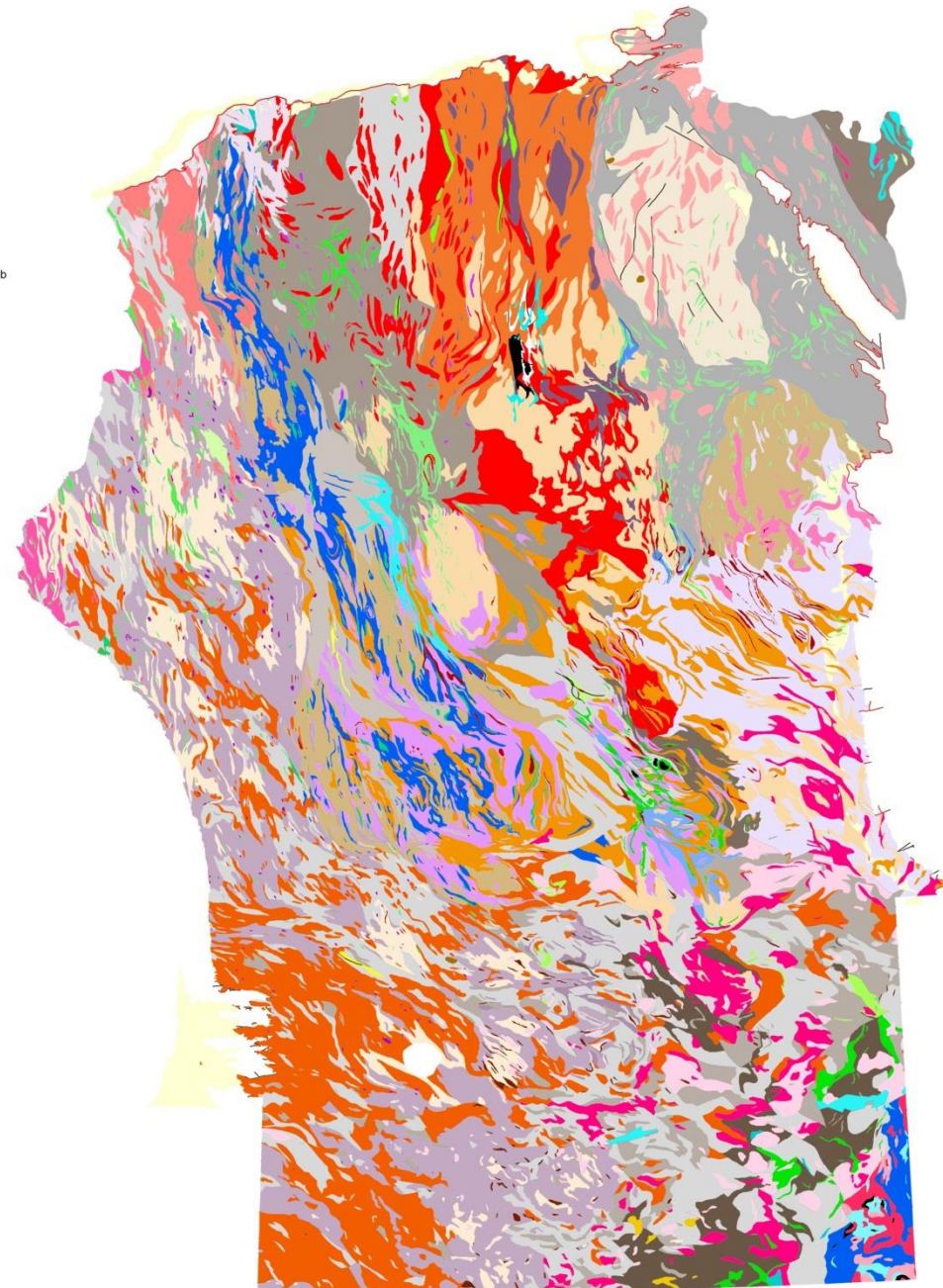


# NEPS: zone fondue

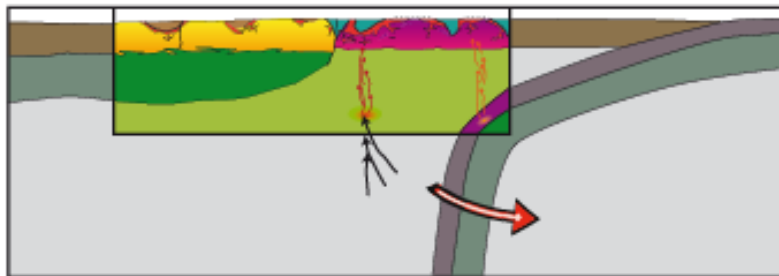
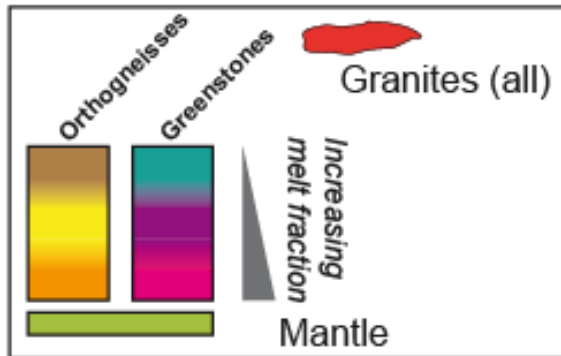
0 20 40 80 120 160  
Kilomètres

## Légende

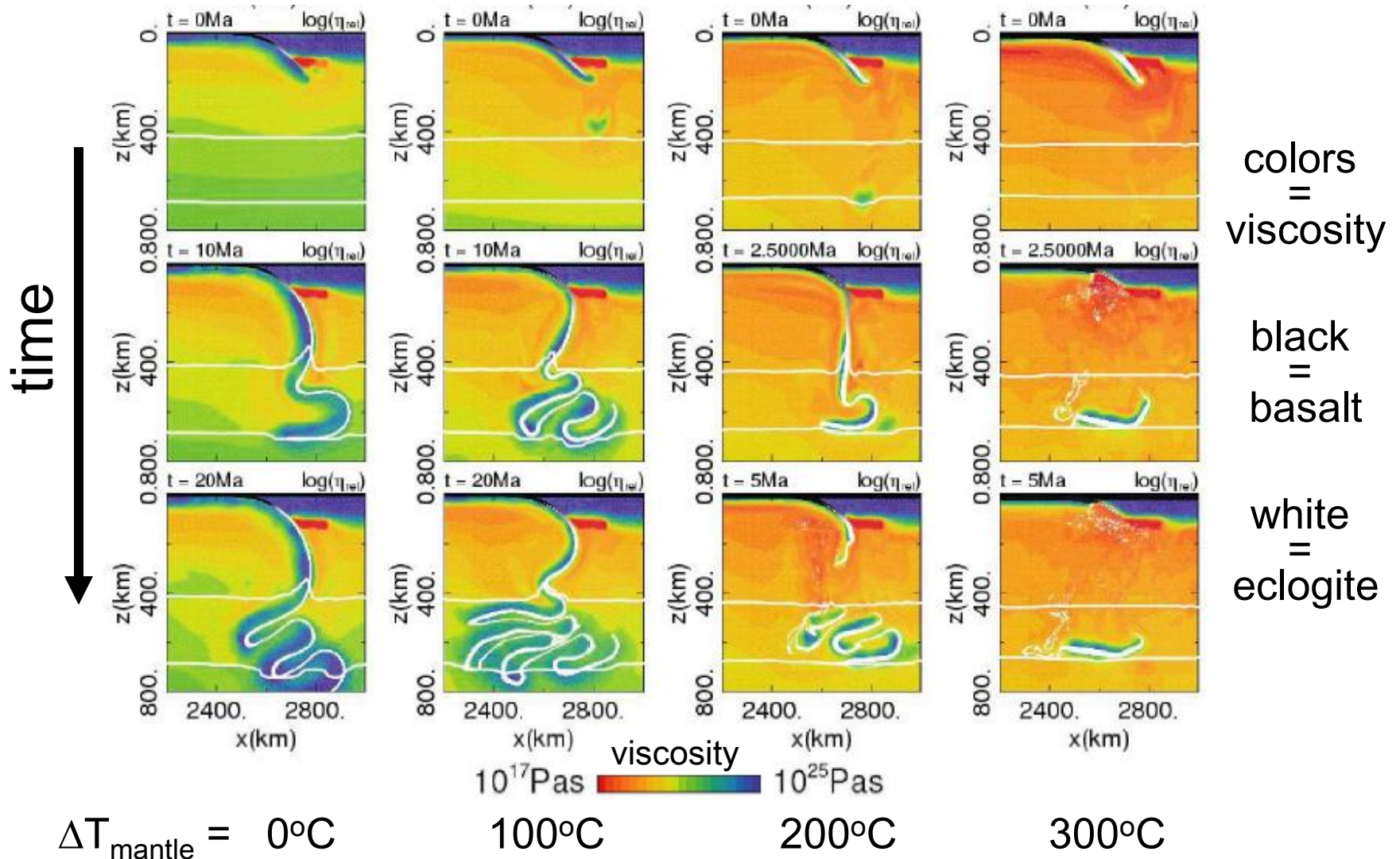
P  
 Ppsa / Akog  
 APdia4  
 APdia3  
 APdia2 / Arlr  
 APdia1 / Apel  
 Ajoy  
 Ader  
 Aopi / Aroy  
 Atst / Abol / Akmb  
 Aram  
 Asan  
 Agdm  
 Acrn  
 Ably  
 Adru  
 Atra  
 Amau  
 Abea / Aqj / Alti  
 Acot  
 Aluk  
 Agrs  
 Aale  
 Ajut / Aat  
 Alma1  
 Alma2  
 Amin / Atie  
 Acmm  
 Alep  
 Aqim  
 Adeb  
 Aduy  
 Arfe  
 Apin  
 Alcv  
 Alrd  
 Alsd / Afav  
 Abyl  
 Abcv  
 Achg  
 Advt / Agat  
 Aviz  
 Achy  
 Atsq / Anan  
 Amel  
 Akkk  
 Asem  
 Accu  
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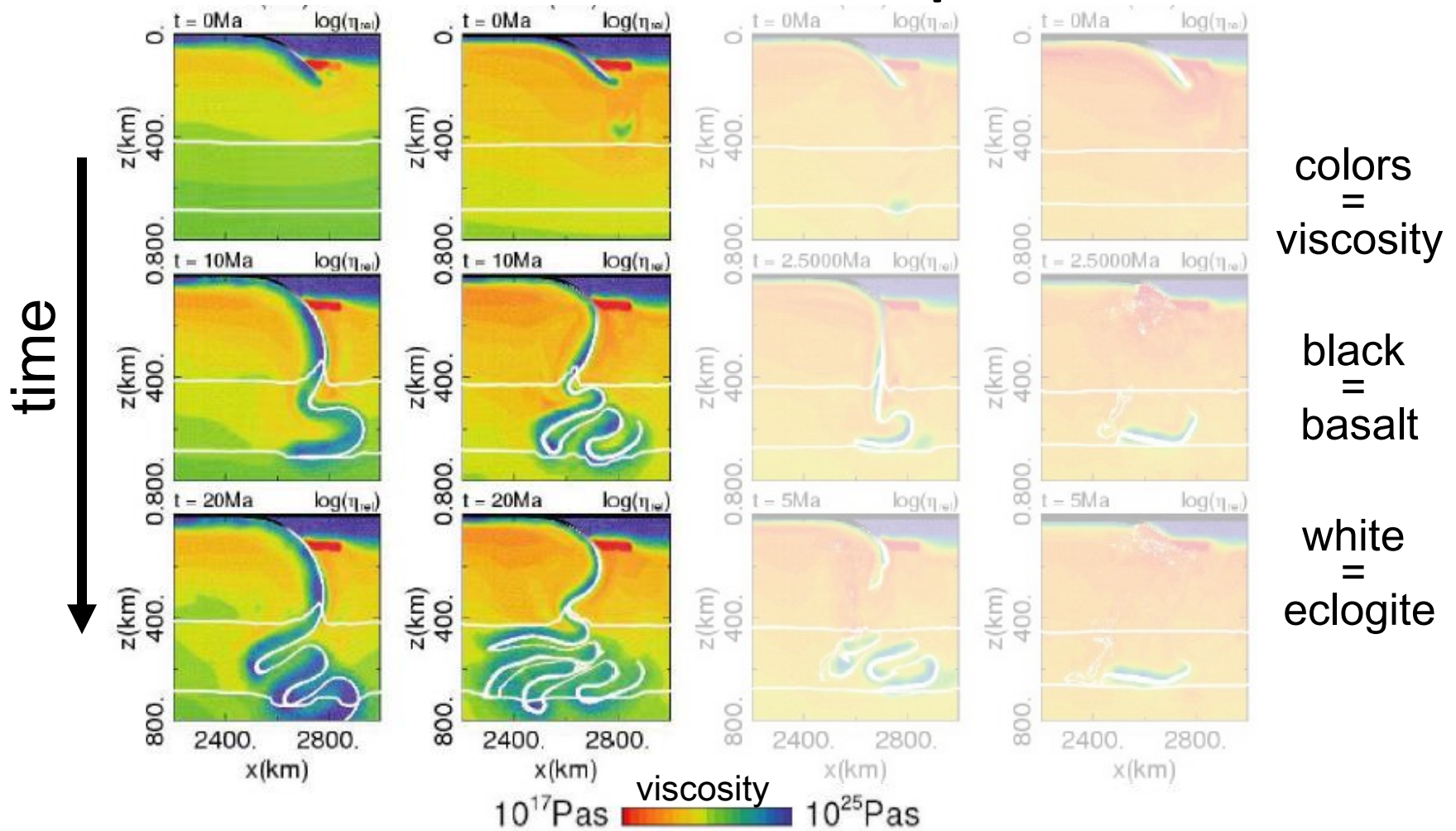
# Considérations géodynamiques



# Subduction ?

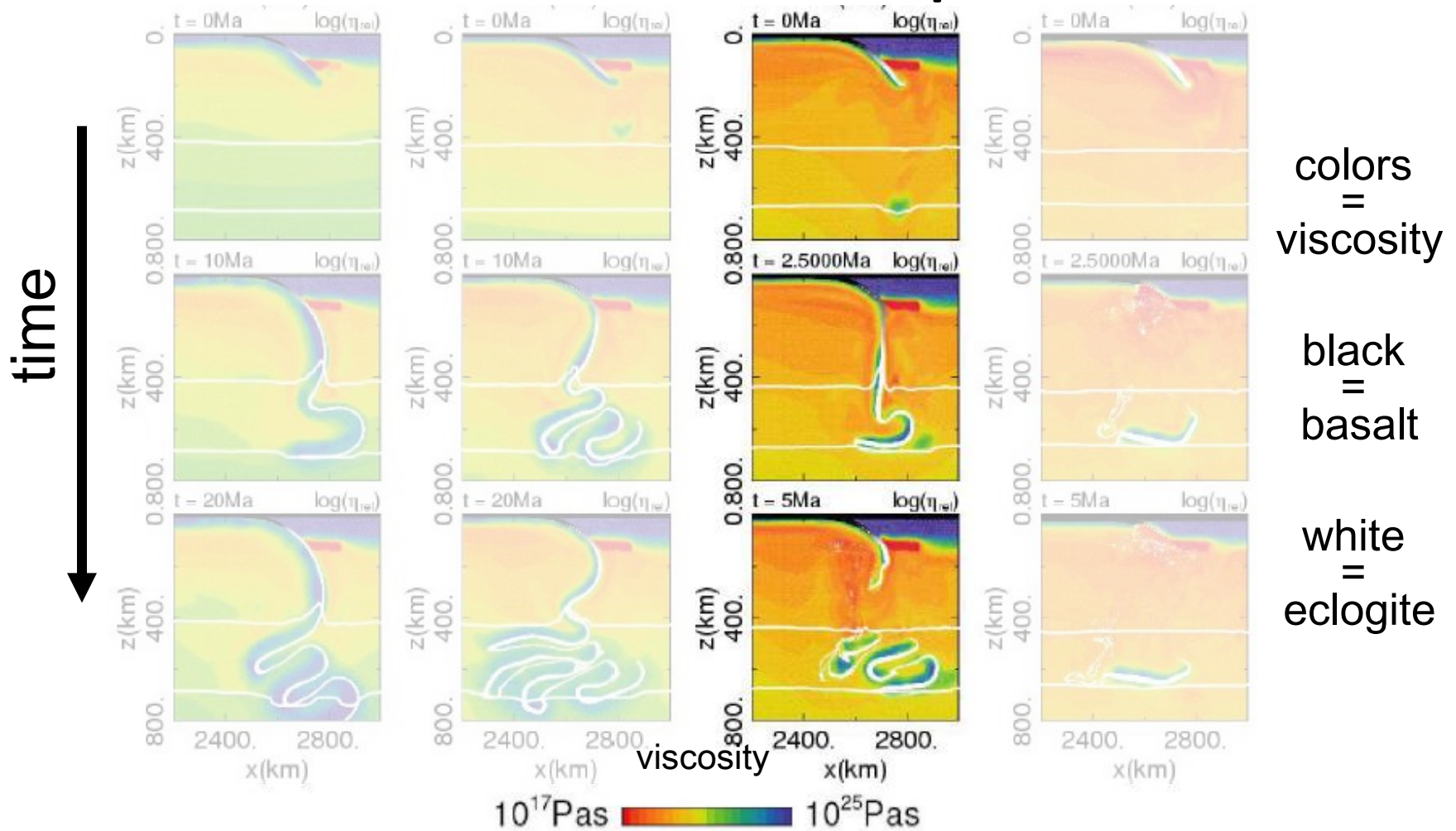


# Weaker Archean plates?



- For low  $T_{\text{mantle}}$  subduction looks like today's

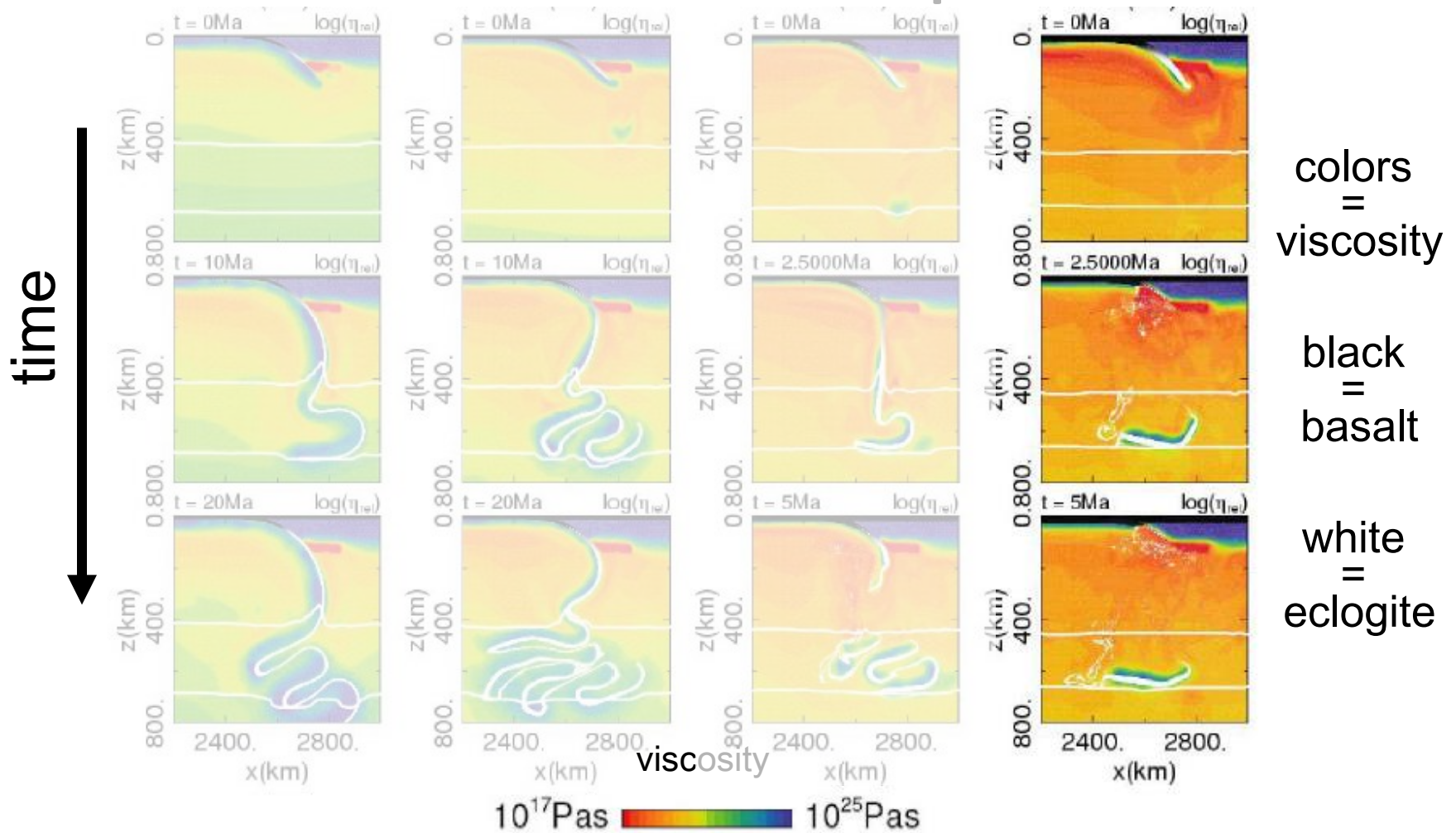
# Weaker Archaean plates?



- For higher  $T_{\text{mantle}}$  frequent slab break-off occurs ...

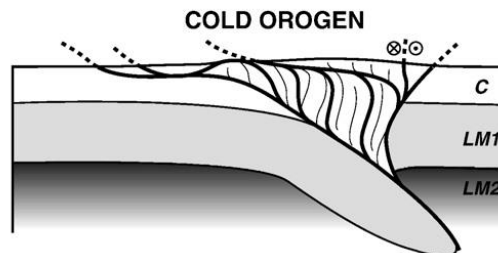
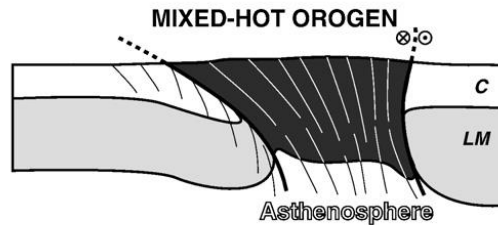
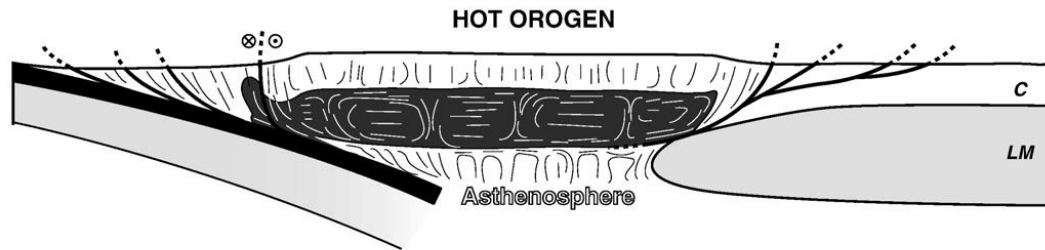
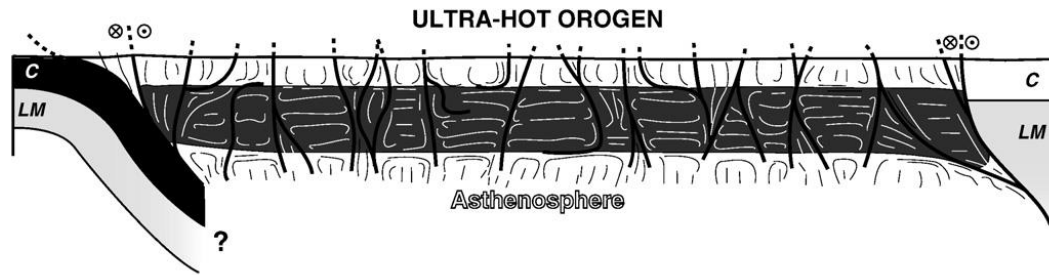
(van Hunen & van den Berg, 2008)

# Weaker Archaean plates?



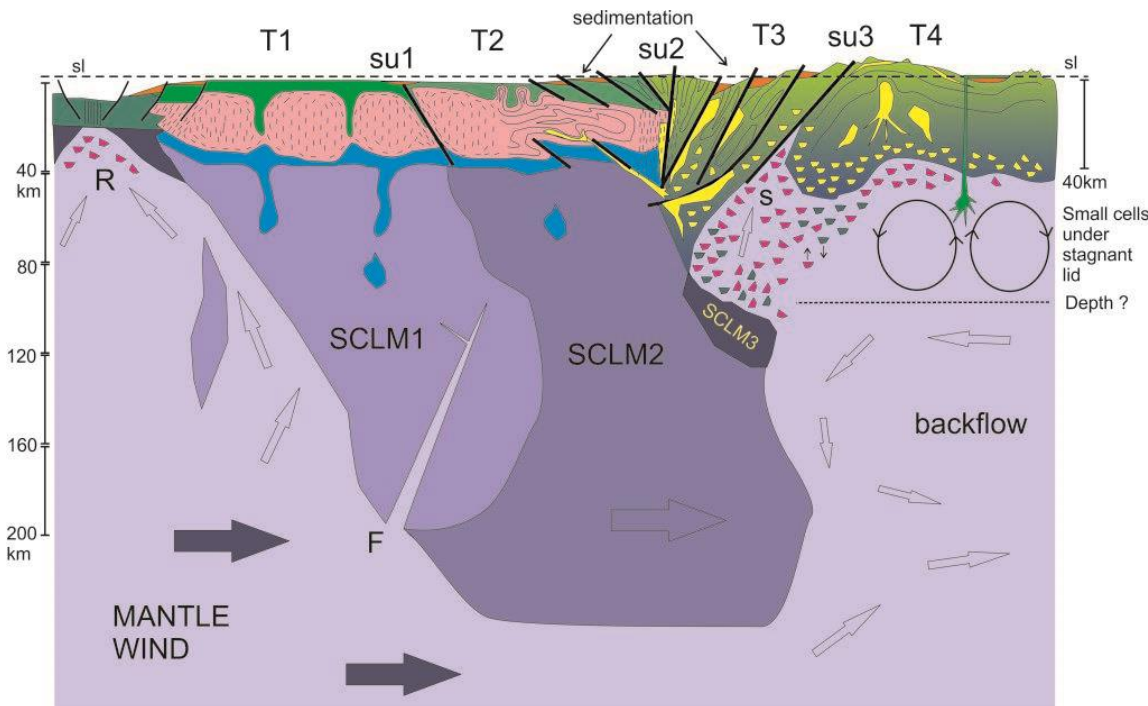
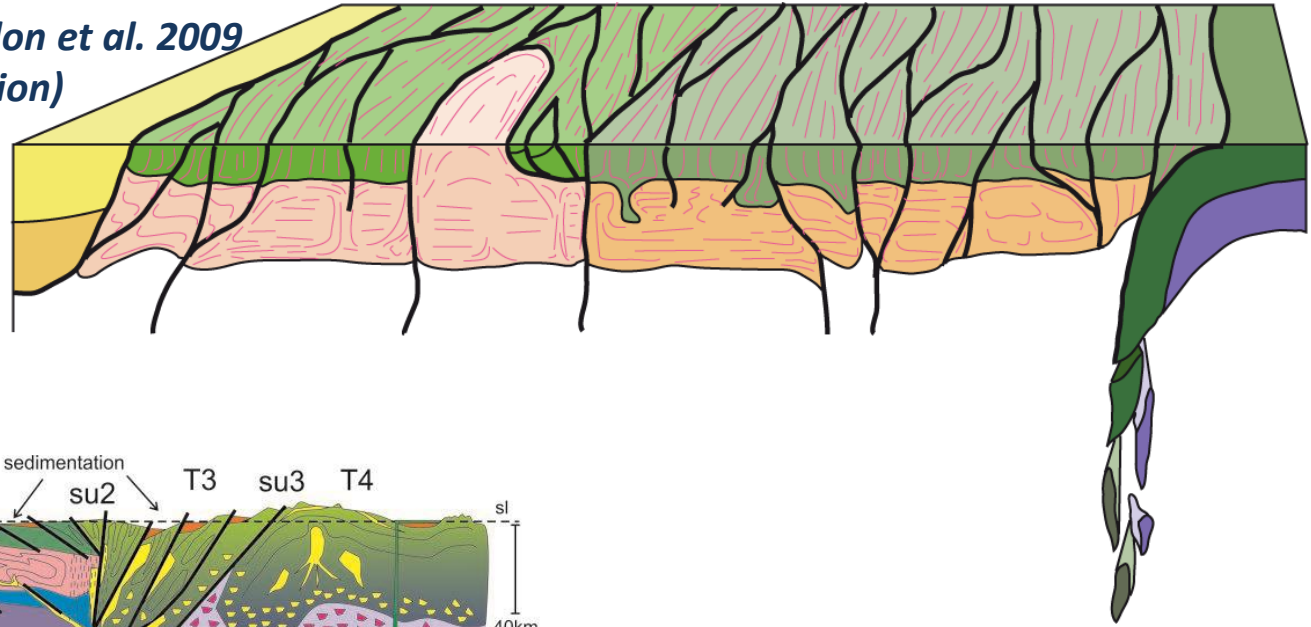
- ... or subduction completely stops.

# Collision ?



# Subduction and collision – they may not look like what we expect

*Hot orogens, inspired by Chardon et al. 2009  
(adding discontinuous subduction)*



*J. Bédard's « snarc-ish » Earth (2012)*

# Le message...

- Les granites reflètent la **construction**, la **différentiation** et la **stabilisation** de la croûte continentale
- Origine **crustale** (surtout) mais il y a aussi des composants **mantelliques**
- Les granites crustaux reflètent surtout leur source, un peu les conditions de fusion: **différentiation progressive de la croûte**
- Les plutons de la croûte supérieure sont les traces de **l'extraction de liquide** de la croûte inférieure fondue

***Merci !***

Barberton (R.S.A.)  
Syénites et granites « tardifs » (3.1 Ga) intrudant les gneiss TTG à 3.45 Ga